Zircon geochronology of bottom rocks in the central Arctic Ocean: analytical results and some geological implications

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ABSTRACT

In the past few years sampling of deepwater seabed gained an increasingly important role in studying geological structure of the Arctic Ocean. A common concept of virtually uninterrupted pelagic drape in the Amerasia Basin and exclusively ice-rafted nature of all clastic components that occur in bottom sediments was challenged by recent discoveries of bedrock exposures in the sea floor, while correlation of results of analytical study of bottom samples collected by the Russian expeditions in 2000, 2005 and 2007 with bathymetric environments at respective sites suggested that certain dredged and cored coarse rock fragments appeared meaningful for bedrock characterization if even the source subpelagic outcrop was not positively documented. The first results of age determinations of detrital zircons that were extracted from coarse fragments of lithic sedimentary rocks resting on the seabed and in the immediate sub-bottom, as well as of zircons from fragments of magmatic/metamorphic rocks and of zircon grains separated directly from sub-pelagic unlithified sediments are in agreement with published interpretations of the Lomonosov Ridge bedrock as composed of Mesozoic terrigenous sequences; the presence of an older Neoproterozoic(?) -Early-Middle Paleozoic basement is also possible. The Mendeleev Rise bedrock, too, is believed to mainly consist of Paleozoic-Early(?) Mesozoic sedimentary superstructure that may locally rest on the Earliest Paleozoic or even older units. Basaltic rocks likely to originate from the High Arctic Large Igneous Province (HALIP) has not so far been found among the collected fragments but limited loose zircon grains probably derived from broadly contemporaneous magmatic products were recorded

in sub-pelagic sediment along with dropstones of variably metamorphosed Precambrian mafic and granitoid rocks.

INTRODUCTION

Great progress in acquisition of new bathymetric and geophysical data relevant to understanding the geological structure and history of the Arctic Ocean, including the tectonic nature of enigmatic Central-Arctic bathymetric highs, was achieved in recent years by the Arctic countries through their programs for delineation of respective extended continental shelves. However, only limited direct geological information was obtained on the composition of sub-bottom bedrock concealed by almost continuous drape of young sediments. Only at a few sites can the lithic fragments recovered by bottom sampling be interpreted with sufficient confidence as representing in situ submarine bedrock, while in most cases they are regarded ice rafted debris (IRD) of questionable derivation.

In search of provenance of lithic and mineral clastic components in bottom sediments we conducted age determinations on zircon crystals of two categories: (1) extracted from the rock fragments and (2) separated directly from hemipelagic sediments. In this paper we present the results of more than 700 zircon U-Pb age measurements completed before 2012. The samples labeled AF00, AF05, AF07 were collected during MS "Akademik Fedorov" cruises Arctic-2000, 2005, 2007, those marked ALR07 were acquired in 2007 on board NIB (nuclear icebreaker) "Rossiya", and two specimens designated BC were selected for the analysis from clastic material sampled by RV "Polarstern" in the course of ARK-XXIII/3-2008 cruise.

Sampling sites were located on Mendeleev Rise, Lomonosov Ridge, on deep Amundsen Basin seabed at the North Pole, and on the bathyal floor in the southern Podvodnikov Basin (Fig 1). Dredging equipment used during Arctic-2000 expedition was supplemented by box and gravity coring on the Arctic-2005 cruise, whereas RV "Polarstern" and the Arctic-2007 cruises employed different types of coring but did not execute any dredging. Sampling on RV "Polarstern" was controlled by Parasound observations which indicated a continuous presence along the ship track of sub-bottom hemipelagic sediments at least several dozen meters thick (Jokat, 2009). Selection of sampling localities surveyed by Russian vessels was only guided by bathymetric data available at the time of cruises.

Zircon dating was performed by high-resolution SIMS method on SHRIMP-II instrument in the Centre of Isotopic Research at VSEGEI, St. Petersburg, Russia. Zircon grains of different morphologies were measured using regular analytical procedure similar to that described by Williams (1998, and references therein) and reference zircons Temora2 (for U/Pb ratios) and 91500 (for U content). Each analytical spot had size ca $2x20x25 \ \mu m$.

DESCRIPTION OF ANALYZED MATERIAL

Zircons in fragments of magmatic and/or metamorphic rocks (Fig. 2)

Almost 200 age determinations, including:

- Station/sample AF07-01 (North Pole): five semi-angular to semi-rounded gravel-pebble size fragments (0.5-0.6 – 1.5-2.0 cm) of granitic rocks with indistinct gneissic banding recovered from box cored pelagic mud. Zircons were analyzed by SIMS SHRIMP directly in thin sections (21 measurements).
- Station ALR07-16: steep western slope of the Geophysicists Spur. Box cored sediments with abundant small rock fragments of variable composition with unusually high proportions of metamorphic and igneous lithologies. Zircon grains were separated from three little splinters of fine-grained gneiss-like rocks and enabled 15 age determinations.





Fig. 1. Location of sampling sites described in this paper. Geological stations designated AF00, AF05 and AF07 were made from MS "Akademik Fedorov" in 2000, 2005 and 2007, respectively, ALR07 from NIB "Rossiya" in 2007, and BC from RV "Polarstern" in 2008. Specimens represented by large fragments and/or pebble-gravel sized debris of zircon-bearing rocks are marked by pentagons (sandstones, siltstones), stars (granitic and gneissic rocks) and squares (metagabbro-dolerites). Triangles indicate samples of hemipelagic sediments. Red lines correspond to the position of small sections of seismic lines shown in Fig. 6.

A single pebble-like fragment of plagiogranite over 2 cm in size from gravity cored sediment (12 U-Pb isotope analyses were made in thin section).

- Stations/samples AF05-08, AF05-24. AF05-26 (dredges) and BC-260 (box corer): southern Mendeleev Rise. Scarce fragments of muscovite, biotite and/or two-mica gneissoid granites and plagiogranites, often cataclastically deformed, gravel-pebble sized, semi-angular to semi-rounded at all sites. One specimen (AF05-08) with distinct gneissic banding had noticeably larger size (8-9 cm) and an almost non-abraded shape. Small pieces of regular petrographic thin sections (without cover glasses) containing visible zircon grains were implanted in standard SIMS mounts (over one hundred measurements).
- Stations/samples AF05-11, AF05-26: southern Mendeleev Rise, dredges. Three small fragments of metagabbro-dolerites among variable other lithologies (25 zircon age determinations in thin sections).

Detrital zircons extracted from fragments of quartz sandstones (Fig. 3)

Stations/samples AF00-05, 10, AF05-11, 14, 15, 20 – different parts of Mendeleev Rise, station/ sample ALR07-18 – Lomonosov Ridge. Numerous sandstone fragments of highly variable size (usually from 1.5-2.0 cm to 10-15 cm, the largest is nearly 40 cm) were recovered by dredges, box and gravity corers and altogether enabled more than 300 zircon age determinations.

Detrital zircons in soft sediments

Station/sample AF07-01 – deepwater seabed at the North Pole, approximately 120 km from the foot of the Lomonosov Ridge. Small portions of soft sediments totaling ~ 300 grams in weight were arbitrary selected from the box cored sample, then mixed and reduced to heavy minerals concentrate which contained about 250 zircon grains. Approximately half of that number appeared unsuitable for age determination (grains too small, or fractured, or filled with inclusions). Unbroken crystals were picked out by hand and analyzed in grain mounts (103 age determinations).



Fig. 2. Morphological appearance of granitoid rock fragments hosting magmatic zircons that were analyzed.



Fig. 3. Examples of the morphological appearance of sedimentary rocks fragments sampled. Arrows indicate the biggest of sandstone fragments dredged at stations AF00-05 & 10 (central Mendeleev Rise) that were selected for age determinations; other debris includes both terrigenous and carbonate rocks some of which are likely to represent IRD. Sandstone specimen ALR07-18 (southern Lomonosov Ridge) was retrieved by gravity core from 55 cm b.s.f.

Station/sample ALR07-15-steep western slope of the Geophysicists Spur 3 km away from the station ALR07-16. A continuous sub-bottom succession was cored to 9 m below sea floor (b.s.f.) and sampled at ~1 m intervals, each sample up to 500 g in weight providing 200-300 small zircon grains. The first 152 measurements reported in this paper were performed on zircons from 12-14 cm b.s.f., 505-507 cm b.s.f. and 703-705 cm b.s.f. (ca 50 grain analyses for each sample).

SUMMARY OF ANALYTICAL RESULTS

The analytical data are presented in the annex (Tables 1 and 2) and illustrated in Figures 4-5. Only concordant or sub-concordant age data were considered for detrital zircons. A brief description of obtained zircon ages is given below.

Fig. 4 demonstrates the lack of apparent correlation between the ages and morphological characteristics of analyzed zircon grains.

Ages of detrital zircons extracted from fragments of sandstones (Fig. 5A):

A common feature of all analyzed specimens is the prevalence of zircons with ages mainly in ~2000 -1000 Ma interval (late Paleoproterozoic – Mesoproterozoic). Yet Precambrian zircons in samples AF00-05 and AF00-10 are mostly late Paleoproterozoic (~2000-1700 Ma), whereas the majority of grains in all other sandstones are Mesoproterozoic (~1800-1700 – 1000 Ma). Another peculiarity of AF00-05 & AF00-10 sandstones is the paucity of Archean zircons relative to the amount observed in other studied sandstones and in soft sediments.

One more distinctive feature of the AF00-05 and AF00-10 specimens is the dominating presence of zircons with Paleozoic to early Mesozoic U-Pb ages whose peaks on the histograms closely resemble the major clusters in hemipelagic sediments. In other sandstones zircons with such ages are absent or very poorly defined.



Fig. 4. Selected SE images of zircon crystals from sandstone specimens AF00-05 and AF00-10 showing lack of correspondence between measured U-Pb ages and the degree of grains roundness.

Ages of zircons in fragments of magmatic and metamorphic rocks (Fig. 5B):

Ages of zircons from granitoid fragments in samples AF07-01, AF05-08, AF05-24, AF05-26 and BC-299 suggest that all listed rocks were mainly crystallized in the Neoarchean (2600 – 2700 Ma). AF07-01 specimens additionally point to the possibility that the parental magma for these granitoids was derived from a Mesoarchean (ca. 2900 Ma) crustal source. Indications of Paleoproterozoic overprint are present in all granitoid samples. The largest and least rounded specimen AF05-08 with the most distinct gneissocity was probably also affected by the Latest Neoproterozoic metamorphic event, as suggested by the presence of rare 600-800 Ma zircon grains with secondary rims.

Granitic rock BC-260 contains only late Paleoproterozoic zircons. The best estimated value of 790±20 Ma obtained on zircons from metagabbro-dolerite specimens AF05-11 and AF05-26 most likely represents the age of magmatic crystallization. Older values close to 2650 Ma and 1950 Ma are closely comparable to ages determined for the granitic rocks and may reflect the presence of zircons captured by mafic magma from older crustal material.

Zircons from three small fragments of gneisslike rocks collected on Geophysicists Spur (ALR07-16) displayed ~1140 Ma, ~570-690 Ma ~400-450 Ma ages. The oldest age was obtained (single shot) on a sole grain recovered from one of the fragments; of three grains extracted from the second fragment two showed ~ 690 Ma (six shots), and one ~ 570 Ma (two shots); and two grains from the third splinter exhibited ~ 407 Ma and ~ 448 Ma ages (three measurements on each grain).







Fig. 5. Distribution of measured zircon ages. A – ages of detrital zircons in sandstone fragments, B – ages of zircons in metamorphosed magmatic rocks, C – ages of detrital zircons in sub-bottom sediments. See text for explanation.

Ages of detrital zircons in Recent sediments (Fig. 5C):

The majority of ages are younger than ~ 500 Ma (Phanerozoic) with lesser peaks in ~ 2000-1800 Ma age interval (late Paleoproterozoic). Neoproterozoc and Meso-Neoarchean determinations are subordinate. A distinct age gap is documented between ~ 1800 and 1000 Ma.

DISCUSSION

The presence of basement outcrops not concealed under sub-pelagic cover or accessible for sampling at shallow sub-bottom depth has been reported, with greater or lesser confidence, from several sites located on the Lomonosov Ridge (Grantz et al., 2001), the southern Northwind Ridge (Grantz et al., 1998), on the central and northern Northwind Ridge, seamounts between Alpha and Northwind Ridges and the southern Alpha Ridge (Andronikov et al., 2008; Brumley et al., 2010, 2011; Database for ECS Dredge Samples at NOAA/ NGDC), and in the central Alpha Ridge (Clark et al., 2000; Jokat, 2003; Van Wagoner et al., 1986). The coarse debris that can positively be attributed to, or inferred to represent the bedrock, is usually mixed with variable proportions of IRD consisting mainly of quartz-rich terrigenous and carbonate rocks. This IRD was defined by Grantz et al. (2011a) as "... shallow marine Paleozoic carbonates and sandstones ... widely distributed on the seabed of the Amerasia Basin by the basin's clockwise Beaufort Gyre current system"; the authors (ibid) further concluded that "...sedimentary clasts in the dredges and cores from Mendeleev Ridge belong to an areally extensive suite of glacial erratics that originated in NW Canada..." Our data suggest that such definition is probably excessively all-embracing, and at least some of the coarse clastic material in sampled bottom sediments on the Mendeleev Rise may appear meaningful for characterization of the local bedrock.

Rock specimens interpreted to represent sub-pelagic basement

The sandstone fragments bearing detrital zircons analyzed in the present study were collected in three different areas – the central Mendeleev Rise, the southern Mendeleev Rise and the near-Siberian segment of the Lomonosov Ridge (see pentagons in Fig. 1). These geographic variations are reflected

in the distribution of detrital zircons ages and other characteristics of respective specimens (Fig. 5A).

The largest of all recovered sandstone fragments were dredged on a small, steep-sided bathymetric spur in the central Mendeleev Rise (sites AF00-05 & 10). Three fragments were analyzed and displayed only slightly differing zircon age data (Fig. 5A, a-c) notably dissimilar to those in the sandstones from the southern Mendeleev Rise. The marked distinctions of these data, such as well expressed Paleozoic-Early Mesozoic zircons population, prevalence of Paleoproterozoic ages over Mesoproterozoic determinations and almost total lack of Archean grains, suggest clastic input from the sources independent from those involved in formation of the sandstones dredged farther south. The central sites are also peculiar for the occurrence of fossiliferous Paleozoic limestones (Kaban'kov et al., 2004) not encountered elsewhere in the sampled area. In our view, these features are likely to signify that AF00-05 & 10 sandstone/carbonate debris represents local Paleozoic and Mesozoic (mostly pre-200 Ma?) sedimentary bedrock strata whose upper horizons may be broadly correlative with sub-pelagic basement of the Lomonosov Ridge described by Grantz et al. (2001) and exemplified in our collection by the specimen ALR07-18 discussed below.

A common feature of specimens from the southern Mendeleev Rise is the predominance of Mesoproterozoic detrital zircons (Fig. 5A, d-h). Sandstones AF05-11[2] and AF05-20 which contain only pre-1000 Ma grains can in reality be as old as Neoproterozoic; this may or may not also be true for the specimen AF05-11[1] where the ~ 200-400 Ma zircon ages are probably too rare to be meaningful. However, more numerous ~ 400-600 Ma grains in specimens AF05-14 & 15 (Fig. 5A, f-g) seem to preclude their Precambrian age; these sandstones also contain lesser amounts of Mesoproterozoic grains and a greater number of ancient grains, some of them as old as Mesoarchean.

Unless caused by the shortage of analytical data, such peculiarities may suggest that sandstones collected at stations 14, 15 and those recovered at stations 11 and 20 differ in age and origin, despite geographical proximity of these sites and apparent lithological similarity of the studied rocks. They

also further confirm the dissimilarity of the southern and the central Mendeleev Rise specimens. If corroborated by subsequent studies, these distinctions would seem easier to explain by local derivation of the analyzed rocks than by their ice rafting from remote sources and selective unloading at different Mendeleev Rise locations. For instance, the presence of Archean grains captured in the analyzed sandstones indicates that these rocks could not be derived from the nearest coastal mainland - the Arctic Alaska-Chukotka (AAC) terrane which was shown by Akinin et al. (2012) to lack the Archean juvenile crust.

The Lomonosov Ridge specimen ALR07-18 is composed of coarse quartzose siltstone with carbonate cement. Zircon U-Pb age data (Fig. 5A, i) indicate input from sources ranging in age from Paleoproterozoic to possibly as young as Early Mesozoic. Except some clustering at about the Paleo/ Mesoproterozoic boundary, the distribution of ages is relatively flat throughout more than a 2000 Ma time interval suggesting multiple recycling of primary clastic material. Lithological composition of the analyzed rock, its likely post-Triassic depositional age and detrital zircons population are consistent with the characterization of the Lomonosov Ridge bedrock by Grantz et al. (2001). The location of sampling site at the base of steep Lomonosov Ridge slope in close vicinity to the near-bottom high of the acoustic basement (Fig. 6a) and a sharply angular shape of the collected specimen suggest possibility of its derivation from a proximal submarine outcrop.

Among magmatic/metamorphic rock fragments the most likely representatives of bedrock were recovered by box corer at the Geophysicists Spur at site ALR07-16 (Fig. 1, Rekant et al., 2012). Here the unusual abundance of fragments is accompanied by uncommonly large amount (about 50%) of metamorphic rocks which at all other sampling sites are invariably markedly subordinate to unaltered carbonate and terrigenous clasts. Increase in overall concentration of coarse material could be caused by slumping of sediments and washing out of fine particles - the processes likely to occur on a steeply faulted slope (Fig. 6b); however, the remarkably high proportion of magmatic/metamorphic rock fragments is uncharacteristic of IRD and, when considered together with bathymetric profile at the sampling site, suggests supply from the local bedrock.

All box-cored rock splinters were too small for preparation of thin sections or chemical treatment. So far only three of them that could visually (using binocular microscope) be defined as gneisses of probable diorite composition were analyzed and showed different (~1140, ~570-690 and~400-450 Ma) ages. In the absence of detailed examination of the mineral composition, metamorphic grade, magmatic vs sedimentary origin, etc. of the samples studied, these ages could be interpreted as indicating that analyzed rocks belong to either the same polymetamorphic assemblage, or are derived from different metamorphic sources. The latter possibility, however, seems highly unlikely, since it would imply transportation of one piece from a Mezoproterozoic provenance, another from a Late Neoproterozoic terrane, and the third from an Early-Middle Paleozoic

area. We therefore prefer the alternative option which allows correlation of the Geophysicists Spur bedrock with the basement assemblages reported from the Northwind Ridge (Brumley et al., 2010, 2011; Database for ECS Dredge Samples at NOAA/ NGDC) and characterized by an ancient (no younger than Grenvillian) protolith affected by subsequent events as young as the Caledonian.

Rock fragments of questionable origin

Interpretation of mineral particles and/or relatively small rock pieces in bottom samples as IRD or otherwise relocated matter (e.g. Bischof et al., 1996; Clark et al., 1980; Grantz et al., 2011a; Phillips and Grantz, 2001) in all probability applies to those subordinate fragments in our collection which are characterized by predominantly small size, subrounded or pebble-like shape and, in some cases, display apparent association with glacial-dominated



Fig. 6. Fragments of stacked seismic sections showing the position of sampling sites ALR07-18 (a) and ALR07-16 (b) relative to the bathymetry and basement behavior (modified from lines shot during Arctica-2011 cruise). Vertical exaggeration approximately 8:1. See Fig. 1 for the location of imaged sections.

layers in sub-bottom sediments. These features are inherent in the majority of analyzed magmatic rocks, namely the granitic pieces AF07-01, AF05-24 & 26, BC-260 & 299 and metagabbro-dolerites AF05-11 & 26 (Figs. 2 and 5B, b-h) which are therefore interpreted as dropstones of questionable origin.

Specimen AF05-08 (Figs. 2 and 5B, i) is distinct among granitoid rocks being much larger with pronounced gneissic banding and almost unsmoothed shape. Its zircon population is characterized by higher amount of Late Archean zircons and the presence of 600-800 Ma grains with secondary rims suggesting Latest Neoproterozoic overprint. The combination of these features may indicate the provenance more proximal than the source area of other Archean granitoids.

Mineral grains interpreted as IRD

Detrital zircons from soft bottom sediments have so far been studied in only two samples collected on different flanks of the Lomonosov Ridge (stations AF07-01 & ALR07-15, Figs. 1 and 5C). Ages of zircons selected at different levels from the ALR07-15 core appeared barely distinguishable (Fig. 5C, b-d). This can either be attributed to invariability of sources that supplied zircon grains to the sampling site during the time spanned by the cored interval, or may merely reflect intermixing of loose sediments on a steep slope. Other findings attracting attention in Fig. 5C are (1) general similarity of zircon age data obtained in radically diverse geographical and geomorphological environments - site AF07-01 in deepwater Amundsen basin (Fig. 5C, a) vs. site ALR07-15 on a prominent bathymetric spur (Fig. 5C, b-d), (2) presence at both localities of post-Triassic grains not recorded in any of the analyzed rock fragments, and (3) notable absence of Mesoproterozoic zircons which constitute the most characteristic population in the studied sandstone specimens.

The first two observations can be interpreted as signifying either a common source or separate but closely comparable provenances. The latter would at first glance seem represented by proximal Paleozoic-Mesozoic sedimentary bedrock reportedly sampled on the Lomonosov Ridge (Grantz et al., 2001) and, based on interpretation of our zircon data from sandstone specimens, also thought to occur on the Mendeleev Rise, at least in the vicinity of sampling sites in the central part. However, upon closer examination such explanation appears difficult to accept. For the North Pole site it would be hard to imagine how abundant heavy mineral products eroded from the Lomonosov Ridge sedimentary bedrock could be delivered to the sampling locality across more than 100 km of flat deepwater Amundsen Basin, and in case of the Geophysicist Spur our data suggest that the bedrock here is more likely composed of Late Precambrian-Early Paleozoic metamorphic basement than of younger rocks capable of releasing post-500 Ma zircons into pelagic sediment.

Ice rafted zircons in sub-pelagic sediments appear therefore the most likely possibility, if even the light minerals and clay components in these deposits could be supplied to both sampling localities by turbidity currents from a variety of sources, including as distal ones as the Laptev Sea shelf. As shown by Krylov et al. (2008) on the basis of ACEX data, in post-Middle Miocene time zircon was a steady component (6-8%) of the heavy minerals assemblage continuously delivered to the Lomonosov Ridge and adjacent bathymetric deeps by Transpolar ice drift from the Arctic margin of Eastern Asia. Consequently, the Phanerozoic and Neoproterozoic zircons could easily be derived from various geological formations of respective age mapped in this extensive region. The provenance of Early Precambrian zircons is more problematic. They could either be supplied from the same enigmatic shield sources which gave rise to the above mentioned magmatic/metamorphic dropstones, or assumed to originate from younger igneous rocks containing inherited ancient grains that were captured by parental melts.

The youngest detrital zircons in Recent subbottom sediments are Late Cretaceous. In all probability they mostly originate from HALIP and/ or broadly contemporaneous volcanic products which are exposed on the Circum-Arctic mainland and islands (Akinin and Miller, 2012; Korago et al., 2010) and believed to extend throughout much of the central Arctic Ocean (e.g. Grantz et al., 2011b). However, based on geophysical data the near-Pole to Russia segment of the Lomonosov Ridge is commonly excluded from the area affected by Late Mesozoic volcanic activity and therefore can hardly serve as a local source for zircons of that age. This further strengthens the notion of their distal derivation and ice rafted nature

The virtual absence in modern deposits of Mesopoterozoic zircons indicates that mineral grains in pelagic sediments were not recycled from sandstones disseminated on the seabed. In case of such recycling the zircon population in sub-bottom layers would be dominated by Precambrian rather than Phanerozoic ages.

CONCLUSIONS

Intensification in recent years of bottom sampling in the central Arctic Ocean was accompanied by implementation of improved methods of site control and state-of-the-art analytical studies of the collected material. This enabled more exact examination of the nature of recovered bottom specimens and expanded the opportunities for interpretation of their lithological and age characteristics in regional geological context.

Our zircon geochronological data suggest that the sandstone/carbonate fragments dredged on the central Mendeleev Rise at sites AF00-05 & 10 most likely represent local Paleozoic and Mesozoic (mainly pre-200 Ma?) sedimentary bedrock units. The youngest of the central Mendeleev Rise sandstones may be broadly correlative with subpelagic Mesozoic sedimentary bedrock of the Lomonosov Ridge confirmed by sampling near the North Pole (Grantz et al., 2001) and believed to be exemplified in the Pole to Siberia segment of the ridge by our specimen ALR07-18 of coarse quartzose siltstone.

The presence of post-500 Ma sandstones among the samples from the southern Mendeleev Rise is more questionable, since the analyzed specimens from this area provided so far only a much lower number of zircons younger than 1000 Ma. At the same time, the well expressed population in these rocks of Mesoproterozoic grains is not a sufficient argument in favor of derivation of the analyzed sandstones from local Neoproterozoic bedrock, as assumed by Kaban'kov et al. (2004, 2008, 2012). While not ruled out by the available data, such possibility requires a much stronger confirmation. Distribution of Precambrian grains in detrital zircon population from Cambrian quartzites in the Canadian Arctic (Hadlari et al., 2012) is very similar to that observed in our specimens of quartzose sandstones from the southern Mendeleev Rise. Consequently, the latter are not necessarily Neoproterozoic and may also be Cambrian or younger, and until their inferred local provenance is constrained with better confidence, the derivation of these rocks from the Canadian provenance and transportation by ice to the Mendeleev Rise will be difficult to disprove.

Zircon geochronology of sandstone debris in bottom sediments from the Mendeleev Rise and the Lomonosov Ridge suggests that these submarine highs are largely underlain by Paleozoic-Early Mesozoic sedimentary bedrock. The latter may in places include Early-Middle Paleozoic fold basement, but the predominance of younger (Middle Paleozoic to Early Mesozoic) platformtype or transitional sequences seems a more likely possibility.

Limited evidence for the presence of older assemblages is provided at the Geophysicists Spur basement high interpreted to consist of metamorphic rocks of possible Grenville-Caledonian affinity.

The source of variably metamorphosed Late Precambrian mafic and Archean granitoid rocks interpreted as dropstones is uncertain. One of many probabilities is that during the glacial maximum they could be scoured by ice from the shallowest blocks of the Lomonosov Ridge some of which may, by analogy with the plateau described by Jackson and Dalh_Jensen et al. (2010), be composed of ancient(?) high-velocity crystalline infrastructure virtually uncovered by sediments.

On the whole, the preliminary geological implications of the present study are consistent with the models proposing significant extension of mature continental crust as a leading mechanism of formation of the Amerasia Basin (Miller et al., 2006; Laverov et al., 2013).

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 Table 1. U-Pb SHRIMP-II analytical data.

Analysis	% 206m	ppm	ppm Th	206 ppm	232-TL /238	(1) 206	Pb/ ²³⁸ U	(1) ²⁰⁷ Pt	o/ ²⁰⁶ Pb	%	(1) 23811/206 DL *	1.0/	(1) 207pt */206pt *	.0/	(1) 207 pt */235 rt	1.0/	(1) 206 _{DL} */238 _L T	.0/	err
number	²⁰⁰ Pb _c	U ts from the	Th North Po	Pb*	nole AE07-01	Age (Ma	a)±abs	Age (Ma)) ± abs	Dis.	U/Pb	±%	Pb Pb	±%	Pb	±%	PbU	±%	corr
79.1	0.00	128	82.3	1.47	0.66	85.2	2.3	286	213	235	75.17	2.7	0.0520	9.3	0.095	9.7	0.0133	2.7	0.28
91.1	0.64	2121	613	30.6	0.30	107	1.0	116	123	9	59.98	1.0	0.0483	5.2	0.111	5.3	0.0167	1.0	0.18
52.1	0.00	276	133	2.24 5.43	0.39	132	2.7	672 128	98	-12	48.29	2.1	0.0619	6.2 4.2	0.177	6.6 4.4	0.0207	2.1 1.4	0.31
57.1	0.45	488	147	9.78	0.31	148	1.6	80	144	-46	43.01	1.1	0.0476	6.1	0.153	6.2	0.0232	1.1	0.17
27.1	0.00	68.6	52.7	1.62	0.79	175	4.1	126	245	-28	36.31	2.4	0.0485	10	0.184	11	0.0275	2.4	0.22
39.1	0.00	220	102	6.40	0.48	214	3.4	202	96	-6	29.57	1.5	0.0493	4.1	0.193	4.4	0.0283	1.5	0.36
1.1	1.38	893	421	27.0	0.49	220	3.0	270	216	22	28.78	1.4	0.0516	9.4	0.247	9.5	0.0347	1.4	0.14
18.1	0.00	33.2 37.0	39.3 58.9	1.00	1.22	222	8.0 6.7	343 284	238	54 23	28.49	3.7	0.0533	11 9.8	0.258	11	0.0351	3.7	0.33
30.1	0.00	323	220	11.0	0.70	250	3.4	231	83	-7	25.32	1.4	0.0508	3.6	0.277	3.9	0.0395	1.4	0.36
9.1	0.00	479	157	16.6	0.34	255	3.4	252	75	-1	24.75	1.4	0.0512	3.2	0.285	3.5	0.0404	1.4	0.39
44.1	1.11	100	177	3.53	1.20	256	5.7	69	303	-73	24.70	2.3	0.0314	13	0.265	13	0.0405	2.3	0.45
45.1	0.00	53.4	143	1.86	2.77	256	6.1	157	184	-39	24.68	2.4	0.0492	7.9	0.275	8.2	0.0405	2.4	0.30
99.1 2.1	0.00	612 35.6	418 67.7	21.5	0.71	258 268	2.6 8.6	288 586	54 233	11 118	24.46 23.53	1.0	0.0521	2.4	0.293	2.6	0.0409	1.0	0.39
93.1	0.58	338	216	12.6	0.66	272	3.3	156	149	-43	23.17	1.2	0.0492	6.4	0.293	6.5	0.0432	1.2	0.19
49.1	1.51	78.2	84.8	2.99	1.12	276	6.8 5.2	403	329	46	22.81	2.5	0.0548	15	0.331	15	0.0438	2.5	0.17
46.1	0.60	172	175	6.52	1.06	278	4.6	228	164	-18	22.72	1.7	0.0507	7.1	0.308	7.3	0.0440	1.7	0.13
75.1	0.00	68.5	73.5	2.62	1.11	281	6.3	566	146	102	22.48	2.3	0.0590	6.7	0.362	7.1	0.0445	2.3	0.32
55.1 42.1	0.83	172 144	405 165	6.66 5.59	2.44	282 283	3.9 4.9	235	157 198	-60 -17	22.33	1.4	0.0483	6.7 8.6	0.298	6.8 8.7	0.0448	1.4 1.8	0.21
51.1	1.01	95.5	110	3.85	1.19	293	5.8	328	245	12	21.52	2.0	0.0530	11	0.339	11	0.0465	2.0	0.19
95.1	0.37	167	162	6.85	1.00	299	4.5	198	170	-34	21.08	1.5	0.0501	7.3	0.327	7.5	0.0474	1.5	0.21
87.1	2.45 0.70	142	62.1 54.1	5.91	0.39	303	5.0	222	189	-144	20.81	2.7 1.7	0.0436	8.2	0.285	8.3	0.0474	2.7 1.7	0.09
89.1	0.39	172	93.7	7.20	0.56	306	4.5	318	164	4	20.60	1.5	0.0527	7.2	0.353	7.4	0.0485	1.5	0.21
21.1 19.1	0.44	410 63.0	367 80.9	17.3	0.93	308 312	4.1 6.9	250 286	122 153	-19 -8	20.44 20.14	1.4 2.3	0.0512	5.3 6.7	0.345	5.5 7.1	0.0489	1.4 2.3	0.25
17.1	0.18	698	228	29.8	0.34	312	3.9	305	67	-2	20.14	1.3	0.0525	2.9	0.359	3.2	0.0497	1.3	0.40
15.1	0.00	372	88.7 56.1	16.0	0.25	316	4.8	287	81	-9 21	19.93	1.6	0.0520	3.5	0.360	3.9	0.0502	1.6	0.40
92.1	0.00	568	437	25.0	0.32	319	3.1	340	73	6	19.74	1.0	0.0543	3.2	0.379	6.2 3.4	0.0507	1.0	0.33
16.1	0.00	252	202	11.2	0.83	325	5.1	370	90	14	19.34	1.6	0.0540	4.0	0.385	4.3	0.0517	1.6	0.37
40.1	0.34	486	162 533	22.0	0.35	330	4.3	308	104	-7	19.04	1.3	0.0525	4.6	0.380	4.8	0.0525	1.3	0.28
38.1	0.68	128	124	5.92	1.00	335	5.6	373	194	11	18.72	1.7	0.0541	8.6	0.398	8.8	0.0534	1.7	0.20
33.1	0.00	387	267	18.0	0.71	339	4.4	294	66	-13	18.51	1.3	0.0522	2.9	0.389	3.2	0.0540	1.3	0.42
76.1 88.1	0.39	614 362	92.1 216	29.5 17.6	0.15	349	3.4	356	63 72	0	17.96	0.9	0.0658	3.0	0.505	3.2 3.3	0.0557	1.0 0.9	0.32
7.1	0.00	206	122	10.0	0.61	355	5.3	362	88	2	17.67	1.5	0.0538	3.9	0.420	4.2	0.0566	1.5	0.37
14.1	0.78	192 1365	92.2 2762	9.47 73.9	0.50	357	5.7	332	194 36	-7	17.54	1.6	0.0531	8.6 1.6	0.417	8.7	0.0570	1.6	0.19
48.1	0.44	213	90.3	11.7	0.44	397	5.8	416	116	5	15.72	1.5	0.0551	5.2	0.483	5.4	0.0636	1.5	0.28
81.1	0.00	86.5	64.1	4.80	0.77	403	8.4	380	117	-6	15.49	2.2	0.0542	5.2	0.483	5.7	0.0646	2.2	0.38
25.1	0.65	262 179	185	15.3	0.73	422 440	6.2	395 436	74	-6	14.80	1.5	0.0546	7.6 3.3	0.509	7.8 3.7	0.0676	1.5	0.19
63.1	0.61	92.6	53.1	6.03	0.59	468	6.4	485	161	4	13.27	1.4	0.0568	7.3	0.591	7.4	0.0754	1.4	0.19
71.1	0.45	279	33.8	18.8	0.13	483	5.5	400	124	-17	12.84	1.2	0.0547	5.6	0.587	5.7	0.0779	1.2	0.21
54.1	0.86	43.2	55.6	3.40	1.33	560	4.5	409	209	-12	11.01	2.1	0.0570	9.5	0.713	9.7	0.0908	2.1	0.21
86.1	0.00	44.1	32.4	3.49	0.76	568	12.3	596	128	5	10.85	2.3	0.0598	5.9	0.760	6.3	0.0921	2.3	0.36
78.1	1.74	372	65.9 308	9.78 33.4	0.61	613	9.7 7.9	382 631	46	-38 -2	9.554	1.7	0.0543 0.0608	2.1	0.746 0.877	2.5	0.0997	1.7	0.14
28.1	0.24	114	28.4	12.9	0.26	796	11	766	75	-4	7.607	1.5	0.0648	3.6	1.173	3.9	0.1314	1.5	0.39
29.1	0.00	46.3	40.6	5.35	0.91	814	14	902	91 77	11	7.430	1.9	0.0691	4.4	1.282	4.8	0.1346	1.9	0.39
94.1	0.00	271	66.8	31.6	0.80	813	8.0	765	39	-0	7.352	1.0	0.0647	1.9	1.204	2.1	0.1348	1.7	0.42
24.1	0.25	244	238	31.2	1.01	892	13	870	56	-2	6.739	1.6	0.0680	2.7	1.392	3.1	0.1484	1.6	0.50
5.1 65.1	0.00	114 231	47.7 37.7	14.9 31.0	0.43	915 936	14 8.8	864 935	63 35	-6 0	6.398	1.6 1.0	0.0679	3.0 1.7	1.428	3.4 2.0	0.1526	1.6 1.0	0.47
80.1	0.00	330	183	72.5	0.57	1469	13	1753	18	19	3.908	1.0	0.1072	1.0	3.784	1.4	0.2559	1.0	0.70
32.1	0.06	343 466	237 99 0	86.4	0.72	1658 1707	19 18	1784	19 15	8	3.409	1.3	0.1091	1.0	4.410 4.581	1.7	0.2933	1.3	0.79
84.1	0.00	333	94.0	89.6	0.22	1755	15	1853	17	6	3.195	1.0	0.1133	0.8	4.888	1.4	0.3129	1.0	0.32
53.1	0.08	185	133	50.6	0.74	1779	15	1852	20	4	3.147	0.9	0.1132	1.1	4.961	1.4	0.3177	0.9	0.65
/3.1 10.1	0.23	82.9 399	54.1 151	23.4	0.42	1830	22 20	1852	34 20	1	3.045	1.4	0.1133 0.1129	1.9 1.1	5.128	2.3	0.3284 0.3295	1.4 1.3	0.58
43.1	0.05	135	52.6	39.0	0.40	1863	23	1873	27	1	2.984	1.4	0.1146	1.5	5.294	2.1	0.3351	1.4	0.69
37.1	0.08	227	112	65.6	0.51	1865	21	1856	19	0	2.981	1.3	0.1135	1.1	5.250	1.7	0.3354	1.3	0.78
100.1	0.00	120	58.5 81.6	44.7	0.48	1870	19	1891	24 22	-1	2.972	1.2	0.1129	1.4	5.370	1.8	0.3366	1.2	0.60
10.2	0.36	82.8	23.8	24.0	0.30	1871	32	1858	39	-1	2.970	2.0	0.1136	2.2	5.274	2.9	0.3367	2.0	0.68
85.1 83.1	0.00	172 1397	229 118	50.1 410	1.37 0.09	1882 1893	18 13	1860 1881	26 9	-1 -1	2.950	1.1 0.8	0.1137 0.1151	1.4 0.5	5.316 5.418	1.8 1.0	0.3390 0.3414	1.1 0.8	0.60 0.84
82.1	0.01	218	238	64.5	1.13	1905	17	1897	19	0	2.908	1.0	0.1161	1.0	5.504	1.5	0.3438	1.0	0.70
77.1	0.05	257	211	76.4	0.85	1913	15	1925	15	1	2.894	0.9	0.1179	0.8	5.619	1.2	0.3455	0.9	0.72
56.1	0.21	218 88.6	43.5 54.2	27.0	0.22	1950	22 19	1930	20 24	-1	2.803	1.5	0.1196	1.1	5.796	1.7	0.35491	1.5	0.76
60.1	0.15	1346	20.6	412	0.02	1963	12	1918	8	-2	2.809	0.7	0.1175	0.5	5.767	0.9	0.3561	0.7	0.84
50.1 90.1	0.19	305 429	53.8 304	93.7 135	0.18	1967 2005	22 15	1968 2000	18 14	0	2.801	1.3 0.9	0.1208	1.0 0.8	5.943 6 183	1.6	0.3568	1.3 0.9	0.79
23.1	0.00	106	48.4	33.4	0.47	2009	27	2015	38	0	2.734	1.5	0.1241	2.1	6.256	2.6	0.3658	1.5	0.59
69.1	0.00	346	148	110	0.44	2032	16	2006	14	-1	2.698	0.9	0.1234	0.8	6.306	1.2	0.3706	0.9	0.76
4.1 66.1	0.27	162	20.3 58.2	20.1 51.7	0.33	2033	30 18	2055	40 19	0	2.698	1./	0.1269 0.1254	2.5 1.1	6.484 6.434	2.8 1.5	0.3707	1.7	0.60
3.1	0.05	280	274	98.4	1.01	2209	27	2174	16	-2	2.447	1.5	0.1357	0.9	7.649	1.7	0.4087	1.5	0.84
77.2	0.08	662	348	248	0.54	2332	16	2366	16	1	2.295	0.8	0.1518	0.9	9.120	1.3	0.4358	0.8	0.66

(1) ⁸U/²⁰⁶Pb (1) ²⁰⁷Pb^{*/206}Pb ppm ²⁰⁶Pb* (1) ²⁰⁶Pb/²³⁸U (1) ²⁰⁷Pb/²⁰⁶Pb % Dis (1)(1)Analysi err ppn ppm Th ⁷Pb^{*/235}U Pb^{*/238}U 232Th/238U +% ge (Ma) ± abs Age (Ma) ±% +% ±% ⁰⁶Pb Ū corr umber abs 0.4553 41.1 0.00 1026 307 401 0.31 2419 24 2376 -2 0 2.196 0.1527 0.4 9.583 0.94 72.1 0.27 85.3 69.1 0.84 2537 27 2529 2.073 1.3 0.1671 1.3 11.11 1.8 0 4824 1.3 0.70 35.4 22 15 35.1 0.09 122 1.72 30 2646 0.84 203 52.5 2609 2.004 1.4 0.1793 0.9 12.33 1.7 0.4988 1.4 20.1 0.05 174 0.30 2634 2740 1.2 0.1897 13.21 0.5048 0.54 401 116 26 30 1.981 1.8 2.2 1.2 22 13 9 61.1 0.08 164 53.2 71.7 0.34 2651 2678 1 1.966 1.0 0.1827 0.8 12.82 1.3 0.5087 1.0 0.78 419 107 191 0.26 19 27 2823 1.889 0.9 0.1996 0.5 14.57 1.0 0.5294 0.84 96.1 0.12 2739 0.9 67.1 0.32 101 15.5 46.5 0.16 2761 2755 18 0 -2 1.870 1.2 0.1915 1.1 14.12 13.94 1.6 1.1 0.5346 1.2 0.74 62.1 0.00 39.0 0.21 20 10 1.863 0.9 0.1884 0.6 0.5368 193 89.2 2770 2728 0.9 0.81 0.2101 0.1853 15.62 13.79 60.2 0.00 151 50.0 70.1 0.34 2780 23 18 2906 12 10 5 -3 1.854 1.0 0.7 1.2 1.0 0.5392 1.0 0.81 59.1 0.00 412 191 0.70 2783 2701 1.852 0.8 0.6 0.5398 0.8 0.79 281 1.6 1.3 1.3 1.1 1.6 1.3 70.1 0.00 46.9 25.7 23.2 0.57 2938 37 2902 21 -1 1.732 0.2095 16.68 2.0 0.5773 0.76 0.15 51.0 34.5 25.3 0.70 2938 31 2936 17 0 1.732 0.2139 17.03 0.5773 0.78 64.1 Spur), sample AL 72.8 1.4 ophysi 0.44 7-15 (83' N, 15 Hemipela edim from tl omor 154 Ridg 12-14 o.s.f. 2.37 88.11 1.9 0.0388 30 0.061 30 1.9 0.06 49.1 360 3.51 -428 10 118 50.1 0.43 695 218 12.3 0.32 131 1.5 5 150 -2525 48 57 1.1 0.0461 6.0 0.131 6.1 0.0206 1.1 0.18 0.0507 0.0215 0.150 1046 0.51 137 226 180 40 46.48 4.2 0.25 46.1 521 19.3 1.4 1.1 4.11.1 174 137 0.0489 0.0495 7.4 5.9 7.5 0.0227 0.0228 31.1 0.58 430 662 8.38 1.59 145 1.6 143 44.08 1.1 0.153 1.1 0.15 -1 17 0.00 129 3.90 0.67 145 43.86 6.1 4.7 9.1 199 2.6 170 1.8 0.156 1.8 0.29 20.1 0.00 385 98.5 7.66 0.26 148 2.3 159 104 43.17 1.6 0.0492 4.5 0.157 0.0232 1.6 0.33 92.3 153 0.0541 0.0233 0.07 8.1 3.13 203 4.19 0.47 148 4.1 375 854 42.96 2.8 38 0.173 38 2.8 26.1 1204 381 0.33 148 1.7 164 68 11 42.94 1.1 0.0493 2.9 0.158 3.1 2.7 0.0233 1.1 0.37 0.14 24.1 32.1 0.00 507 139 13.9 0.28 202 2.4 205 56 1 31.38 1.2 0.0502 2.4 0.221 0.0319 1.2 0.45 33.1 172 86.6 4.97 0.52 214 2.9 518 174 60 29.64 1.4 0.0577 7.9 0.268 8.1 0.0337 1.4 0.17 2.1 1.14 160 109 5.36 0.70 244 4.8 189 303 -22 25.92 2.0 0.0499 13 0.265 13 13 0.0386 2.0 0.15 25.77 15.1 12.6 0.52 245 3.8 246 300 0 1.6 0.0511 13 0.273 0.0388 1.6 0.12 1.08 373 187 17.1 13.1 1.26 0.45 255 1.99 131 159 4.60 253 5.1 204 -19 25.00 2.1 0.0502 11 0.277 11 4.4 0.0400 2.1 0.18 1.0 0.0508 0.287 1.0 0.24 1566 675 55.5 259 98 -11 0.0410 0.66 2.7 231 24.39 4.3 11.1 25.1 42.8 542 1.15 1.31 7.4 3.5 194 151 28 -14 8.6 6.5 9.0 6.7 0.00 47.8 1.63 280 357 22.52 2.7 1.3 0.0537 0.329 0.0444 2.7 1.3 0.30 686 22.45 0.0445 0.84 281 242 0.0510 0.313 0.19 20.9 27.1 29.1 398 263 331 134 0.0547 0.0515 0.339 0.322 1.40 200 399 7.84 2.06 284 5.1 40 22.20 1.8 1.2 15 15 0.0450 0.12 1.8 1.2 0.0453 286 0.88 605 520 23.8 0.89 22.07 3.4 -8 4 5.8 6.0 0.20 38.1 0.00 527 17.1 327 20.6 0.64 1.69 287 3.0 15.1 299 44 147 21.96 1.1 0.0523 1.9 0.328 0.825 2.2 9.4 0.0455 1.1 0.48 0.00 329 1868 468 19.09 0.1143 8.2 3.8 0.0524 0.50 3.1 28.1 0.77 4.7 4.7 21.1 0.50 525 526 23.9 1.03 331 3.9 520 83 89 57 23 18.95 1.2 0.0577 0.420 4.0 0.0528 1.2 0.31 1.1 7.1 0.43 536 682 26.2 1.32 355 4.2 438 17.67 1.2 0.0556 4.0 0.434 4.2 0.0566 1.2 0.29 121 570 5.5 25 0.00 75.6 117 3.78 1.61 365 7.2 495 36 17.18 2.0 0.0571 0.458 5.8 0.0582 2.0 0.35 42.1 71.4 0.56 365 8.8 330 -11 17.16 2.5 0.0530 25 0.0583 1.54 38.7 3.57 0.426 2.5 0.10 30.1 0.37 283 277 15.2 1.01 389 5.2 409 102 16.07 1.4 0.0549 4.5 0.471 4.7 0.0622 1.4 0.29 5.0 7.4 -20 10.1 0.68 318 410 17.8 1.33 403 322 151 15.50 1.3 1.9 0.0528 6.7 0.470 6.8 0.0645 1.3 0.19 5.1 2.50 181 95.2 10.5 0.54 343 413 -17 15.13 0.0533 18 0.486 18 0.0660 1.9 0.10 412 16.1 34.1 0.91 174 170 11.5 27.0 1.01 475 7.9 594 270 25 -24 13.08 1.7 1.1 0.0597 12 4.1 0.629 13 4.3 0.0764 1.7 0.14 0.29 410 183 0.46 476 5.0 386 93 13.06 0.0544 0.574 0.0766 1.1 0.26 22.1 1.41 130 62.1 90.7 8.94 0.49 490 8.4 479 287 80 -2 -7 12.66 1.8 1.2 0.0567 13 3.6 0.617 13 3.8 0.0789 1.8 1.2 0.14 14.1 0.36 0.18 460 12.55 0.0562 0.0796 512 35.2 494 5.6 0.617 0.31 41.1 12.1 317 248 1.03 1.23 25 -12 1.2 1.4 2.6 9.5 2.8 9.6 1.2 1.4 0.42 0.15 317 233 529 5.9 693 55 11.70 0.0625 0.737 0.0855 595 209 0.0578 0.770 1.22 8.2 10.34 0.0967 20.8 521 296 40.1 47.1 73.6 146 0.18 0.49 673 774 0.0926 0.0649 1.4 1.4 1.405 1.141 0.93 0.64 0.19 414 24 1480 27 30 57 -0 9.088 3.8 1.2 4.0 0.1100 3.8 1.2 39.1 0.00 307 33.6 8.6 7.843 1.9 0.1275 771 926 195 4.1 0.32 150 224 0.17 1591 1919 21 3.570 1.0 0.1176 0.9 4.537 1.3 0.2799 0.76 14 16 28 15 19 1.0 1.2 4.924 0.3160 28.1 0.21 122 52.9 0.65 1770 19 1848 4 3.163 0.1130 1.5 2.01.2 0.63 35.1 39.1 0.03 239 81.3 65.4 0.35 0.49 1786 18 20 1831 3 4 3.133 1.2 0.1120 0.8 4.927 1.4 0.3192 1.2 1.3 0.82 179 1805 84.7 49.7 1866 3.096 1.3 0.1141 1.0 5.083 1.6 0.3230 0.78 24.1 0.00 222 158 61.9 0.74 1813 19 1873 20 24 3 3.079 1.2 0.1146 1.1 5.130 1.6 0.3247 1.2 0.74 37.1 0.01 639 156 178 0.25 1815 17 1794 -1 3 3.075 1.0 0.1097 1.3 4.918 1.7 0.3252 1.0 0.62 12 27 13 19.1 0.14 1046 237 295 0.23 1828 16 1878 3.049 1.0 0.1149 0.7 5.193 1.2 0.3279 1.0 0.82 18.1 0.13 139 93.0 39.6 0.69 1843 22 1859 1 3.021 1.4 0.1137 1.5 5.188 2.1 1.4 0.3309 1.4 0.68 43.1 259 153 0.61 19 2 3.014 1.2 0.1144 0.7 5.235 0.3317 1.2 73.8 1847 1871 0.85 36.1 23.1 0.03 882 68.3 255 0.08 1872 16 19 1901 52 19 2 2.968 1.0 0.1163 2.9 1.0 5.404 3.1 1.6 0.3369 1.0 0.33 255 134 0.54 -1 2.936 1.2 0.1145 5.378 0.3406 1.2 0.75 0.00 74.7 1890 1872 0.77 0.17 1.3 1.3 1.6 1.6 1.3 1.3 48.1 0.00 192 143 56.7 1902 21 1883 15 15 -1 7 2.913 0.1152 0.9 5.453 0.3432 0.83 44.1 0.00 159 26.5 2237 25 0.1537 0.9 8.790 0.4149 0.83 56.6 2387 2.410 6.1 0.08 681 420 276 0.64 2491 21 2645 10 6 2.119 1.00.1791 0.6 11.65 1.2 1.9 0.4718 1.0 0.87 45.1 270 0.34 1.710 0.2154 0.5849 0.56 537 175 2969 26 2946 1.1 1.6 17.37 1.1 nipel Ridg ophysio Spur) nple AL -15 (83 N, 15 ۱b.s.f 10.1 0.30 1579 29.1 0.03 137 2.9 263 81 93 46.70 2.2 0.0515 3.5 0.152 4.1 0.0214 2.2 0.52 44.837.1 0.41 2.1 36 -73 46.45 1.5 7.4 0.139 7.5 1.5 0.20 0.87 570 226 10.6 137 180 0.0468 0.0215 -52 160 2.2 6.9 0.31 18.1 0.74 1009 273 19.4 0.28 142 3.0 -137 45.04 0.0451 6.6 0.138 0.0222 2.2 0.40 814 2.4 0.0662 4.1 0.247 4.8 0.0270 2.4 0.50 26.1 0.00 222 86.2 5.15 172 4.1 86 373 36.99 920 23.1 3.86 114 60.7 2.81 0.55 176 5.6 65 -63 36.10 3.2 0.0470 39 0.181 39 0.0277 3.2 0.08 0.13 176 2.5 154 -13 1.5 0.0491 2.6 3.0 0.0277 43.1 0.00 1300 169 31.0 60 36.08 0.188 1.5 0.49 2.4 2.4 4.1 1.30 374 684 9.21 1.89 180 4.2 -17 360 -109 35.33 0.0457 15 0.178 15 0.0283 0.16 2.1 2.61 177 371 4.50 2.17 183 4.9 -100 560 34.66 2.7 0.0440 23 0.176 23 0.0289 2.7 0.12 -155 8.1 0.42 822 1087 29.8 1.37 266 5.6 5.6 310 98 17 23.75 2.2 2.1 0.0526 4.3 0.305 4.8 0.0421 2.2 2.1 0.45 27.1 0.37 0.38 266 258 100 -3 23.70 0.0514 4.4 0.299 4.9 0.0422 0.44 703 260 25.6 1.26 0.66 6.4 5.8 2.3 2.1 8.5 21 8.9 21 30.1 0.44 225 275 8.65 281 152 200 -46 22.47 0.0491 0.301 0.0445 2.3 0.26 35.1 124 500 0.0480 0.0446 1.91 79.8 4.85 281 120 -57 22.44 0.297 2.1 0.10 14.1 15.1 125 712 0.87 0.76 7.7 6.4 120 140 2.8 2.2 6.6 6.3 0.00 105 4.78 282 1236 339 22.38 0.0816 6.0 0.503 0.0447 2.8 0.42 0.0494 5.9 0.35 0.66 524 28.8 295 164 -44 21.37 0.318 0.0468 2.2 2.48 0.40 69 760 1.4 3.0 0.0521 0.0440 31.1 0.24 631 1516 25.8 300 4.1 290 -3 21.03 3.0 0.342 3.3 0.0476 1.4 0.42 0.295 4.02 303 -128 0.0481 0.10 3.1 2.68 94.5 36.5 9.0 -84 31 31 20.77 28.1 5.1 442 361 155 0.84 304 7.1 9.7 125 233 400 200 -59 -47 20.70 2.4 2.3 0.0485 0.0508 17 0.323 17 9.0 0.0483 0.0700 2.4 2.3 0.14 4.17 19.1 15.8 0.62 14.27 0.491 0.25 1.11 260 436 8.7 47.1 254 250 1.02 445 6.9 407 14.00 0.0549 2.9 0.541 3.4 0.0714 1.6 0.48 0.00 15.6 66 1.6 34.1 1.19 105 87.1 7.02 0.85 475 8.6 383 330 -19 13.06 1.9 0.0543 15 0.573 15 0.0765 1.9 0.13 1.9 1.6 1.5 46.1 0.24 296 250 152 0.53 492 7.5 547 61 72 11 0.0585 2.8 0.639 3.2 0.0793 0.49 20.2 12.61 1.6 7.3 40.1 0.30 41.4 17.8 0.17 512 518 12.10 0.0577 3.3 0.657 3.6 0.0826 1.5 0.41 0.64 97 1.5 0.0592 0.754 4.7 0.0924 44.1 1.54 945 589 76.2 570 7.9 574 10.81 4.5 0.31 1.5 240 29.1 1 24 120 128 10.5 1.10 614 14 308 -50 10.00 2.5 0.0525 11 0 724 11 0.1000 25 0.23 42.1 1.13 45.3 21.7 4.20 0.50 653 14 575 270 -12 9.380 2.2 0.0592 12 0.870 13 0.1066 2.2 0.17 113 329 2.2 2.2 4.7 2.3 2.2 2.2 17.1 0.68 199 20.4 0.59 722 15 680 100 -6 -1 8 4 3 0 0.0622 1.017 5.2 3.2 0 1 1 8 6 0.43 0.85 48 0.0645 11.1 0.25 402 43.9 769 16 759 7.890 1.127 0.69 0.1267 39.1 25.1 0.77 0.37 1.5 2.2 1.8 2.2 1.219 2.017 0.1322 0.1328 0.00 167 124 18.9 800 11 835 37 40 4 7.570 0.0669 2.3 3.1 1.5 0.64 0.56 271 96.7 31.2 804 17 7.530 0.1102 2.2 0.70 1802 124

Table 1. Continued.

0.16

209

148

23.9

0.73

805

11

753

62

7.520

1.5

0.0643

2.9

1.180

3.3

0.1330

1.5 0.45

36.1

Table 1. Continued.

Analysis number	% ²⁰⁶ Ph	ppm U	ppm Th	ppm ²⁰⁶ Pb*	²³² Th/ ²³⁸ U	(1) ²⁰⁶ I Age (Ma	$(2b)^{238}U$	(1) ²⁰⁷ Pl Age (Ma	$b/^{206}$ Pb	% Dis.	(1) ²³⁸ U/ ²⁰⁶ Pb [*]	+%	(1) ²⁰⁷ Pb ^{*/206} Pb [*]	+%	(1) ²⁰⁷ Pb ^{*/235} U	+%	(1) ²⁰⁶ Pb ^{*/238} U	+%	err
41.1	0.15	191	242	22.6	1.31	831	12	792	40	-5	7.270	1.5	0.0655	1.9	1.243	2.4	0.1376	1.5	0.61
6.1	0.78	77.0	6.28	14.2	0.08	1246	28	1754	69 32	41	4.690	2.4	0.1073	3.8	3.150	4.5	0.2133	2.4	0.54
20.1	0.11	564	314	140	0.48	1636	31	1908	100	17	3.460	2.2	0.1168	5.6	4.650	6.0	0.2889	2.2	0.36
24.1	0.30	275	162	77.0	0.61	1815	34	1853	25	2	3.073	2.1	0.1133	1.4	5.080	2.6	0.3252	2.1	0.84
12.1	0.15	516 100	198 53.5	145 28.8	0.40	1827	34 27	1845	16 31	-4	3.050	2.1	0.1128 0.1093	0.9	5.040	2.3	0.3277	2.1	0.92
32.1	0.13	156	84.3	49.6	0.56	2029	25	1995	19	-2	2.702	1.4	0.1226	1.1	6.250	1.8	0.3699	1.4	0.80
1.1	0.10	551	305	177	0.57	2043	36	1993	12	-2	2.681	2.1	0.1225	0.7	6.300	2.2	0.3729	2.1	0.95
22.1	0.00	223	89.4	91.2	0.30	2522	45	2678	13	-5	2.088	2.2	0.1230	0.8	12.06	2.4	0.3870	2.2	0.92
38.1	0.02	529	203	218	0.40	2529	29	2605	23	3	2.081	1.4	0.1748	1.4	11.58	2.0	0.4805	1.4	0.71
9.1 7.1	0.35	154	95.4 101	63.8 76.4	0.64	2532	48 50	2641 2654	29 14	4	2.076	2.3	0.1787	0.9	11.85	2.9	0.4810	2.3	0.80
16.1	0.01	285	159	126	0.58	2665	46	2682	10	1	1.953	2.1	0.1832	0.6	12.93	2.2	0.5120	2.1	0.96
13.1	0.03	444	372	196	0.86	2672	47	2724	9.3	2	1.947	2.1	0.1879	0.6	13.30	2.2	0.5140	2.1	0.97
21.1	0.00	187	127	84.3	0.56	2704	48	2093	10	0	1.919	2.2	0.1844 0.1879	0.7	13.25	2.3	0.5212	2.2	0.86
49.1	0.00	226	235	103	1.08	2752	33	2727	11	-1	1.878	1.5	0.1882	0.7	13.82	1.6	0.5326	1.5	0.92
45.1 Hemipelagio	0.05 sediment	336 ts from the	80.0 Lomono:	164 sov Ridge	0.25 (Geophysicis	2896 ts Spur), si	34 ample ALR	2800 07-15 (83°	9.2 ° N, 156°	-3 E), 703-70	1.763 5 cm b.s.f.	1.4	0.1969	0.6	15.40	1.6	0.5672	1.4	0.93
18.2	0.47	832	473	9.17	0.59	82	1.5	236	190	189	78.30	1.9	0.0509	8.1	0.090	8.3	0.0128	1.9	0.23
18.1	1.60 0.45	636 1078	503 802	7.22	0.82	83 89	1.7	-15	310 190	-4 -117	76.90 72.00	2.0	0.0476	13 8.0	0.085	13 8.1	0.0130	2.0 1.8	0.15
19.1	0.73	281	287	4.69	1.06	123	2.5	155	200	26	51.80	2.1	0.0491	8.4	0.131	8.7	0.0193	2.1	0.24
26.1	2.02	513	101	9.01	0.20	128	2.6	258	310	102	50.00	2.0	0.0514	14	0.142	14	0.0200	2.0	0.15
23.1	0.28	732	244 280	14.0	0.14 0.40	138	2.5	113	120	-20	46.11 45.08	1.7	0.0491 0.0483	5.2 5.3	0.147 0.148	5.6	0.0217	1.7	0.46
33.1	0.40	521	141	10.2	0.28	145	1.5	123	130	-15	43.95	1.1	0.0485	5.4	0.152	5.5	0.0228	1.1	0.19
20.1	1.56 2.69	289 243	83.3 139	5.80	0.30	147 165	3.1	311	330 510	-92	43.45	2.2	0.0526	15 21	0.167	15 21	0.0230	2.2 2.2	0.15
16.1	0.00	252	113	6.20	0.46	182	3.7	214	99	17	34.85	2.1	0.0504	4.3	0.199	4.8	0.0287	2.1	0.44
14.1	1.67	205	62.5	5.41	0.31	192	4.2	-62	450	-133	33.14	2.2	0.0449	18	0.187	19	0.0302	2.2	0.12
41.1	1.41	508	318	17.4	0.65	240	2.6	295	250	-278	25.46	1.1	0.0390	11	0.208	11	0.0390	1.1	0.10
48.1	4.08	73.1	146	2.58	2.06	249	7.9	459	800	84	25.38	3.2	0.0560	36	0.300	36	0.0394	3.2	0.09
21.1	1.02	247 60.2	408 53 3	8.78	1.71	259 263	5.0 9.2	92 623	250 710	-65 137	24.37 24.00	2.0	0.0479	11	0.271	11	0.0410	2.0	0.18
44.1	0.60	165	179	5.99	1.12	265	3.8	371	160	40	23.79	1.5	0.0540	6.9	0.313	7.1	0.0420	1.5	0.21
32.1	0.26	940	548	34.5	0.60	269	1.9	267	57	-1	23.43	0.7	0.0516	2.5	0.304	2.6	0.0427	0.7	0.28
39.1	0.49	233 447	588	8.78	1.21	272 280	5.5 2.6	-30 275	430 120	-111	23.16	2.1	0.0455 0.0518	18 5.1	0.271 0.317	5.1	0.0432	2.1	0.12
37.1	4.54	88.2	65.3	3.63	0.76	288	7.8	1183	600	311	21.90	2.8	0.0790	30	0.500	31	0.0456	2.8	0.09
11.1	1.07	133	177	5.36	1.38	293	6.1 5.4	322	250	10	21.49	2.1	0.0528	11	0.339	11	0.0465	2.1	0.19
38.1	0.91	486	212	20.7	0.45	300	2.9	378	200	26	21.02	1.0	0.0522	8.7	0.355	8.8	0.0475	1.0	0.11
28.1	0.19	937	220	47.9	0.24	372	6.0	321	52	-14	16.85	1.6	0.0528	2.3	0.432	2.8	0.0594	1.6	0.59
3.1	0.19	875 906	1567	49.7 54.6	1.85	412	6.5 6.9	425	44 58	-11	15.17	1.6	0.0553	2.0	0.503	2.6	0.0659	1.6	0.64
24.1	0.54	218	120	13.4	0.57	442	7.8	445	120	1	14.09	1.8	0.0558	5.2	0.546	5.5	0.0710	1.8	0.33
50.1	2.76	223	112	14.9	0.52	469	6.2	605	250	29	13.24	1.4	0.0600	12	0.625	12	0.0755	1.4	0.12
40.1 49.1	0.20	48.0	414	4.55	0.39	705	5.1	702	45	0	8.654	0.8	0.0628	2.1	1.001	2.2	0.1050	0.8	0.20
6.1	0.05	273	94.6	62.8	0.36	1529	24	1804	18	18	3.736	1.7	0.1103	1.0	4.070	2.0	0.2677	1.7	0.87
25.1	0.07	214	91.8 66.5	56.0 63.1	0.44	1713	25 12	1703	20	-1	3.284	1.7	0.1043	1.1	4.381	2.0	0.3045	0.8	0.84
13.1	0.13	188	121	53.0	0.67	1824	27	1843	21	1	3.057	1.7	0.1127	1.2	5.080	2.0	0.3271	1.7	0.83
15.1	0.12	167	75.0	47.4	0.46	1832	27	1850	20	1	3.041	1.7	0.1131	1.1	5.130	2.0	0.3287	1.7	0.84
10.1	0.06	202	184	58.8	0.94	1877	27	1871	18	0	2.959	1.7	0.1144	1.0	5.330	1.9	0.3379	1.7	0.86
47.1	0.06	520	101	156	0.20	1926	11	1919	13	0	2.872	0.7	0.1175	0.7	5.642	1.0	0.3481	0.7	0.69
43.1 27.1	0.19	642 76.1	596 19.2	217 26.3	0.96	2132 2169	14 35	2773	8.6 34	2	2.551 2.498	0.8	0.1936	2.0	7.670	2.7	0.3919	0.8	0.82
35.1	0.06	231	146	87.4	0.65	2356	17	2366	16	0	2.267	0.9	0.1518	0.9	9.230	1.3	0.4411	0.9	0.69
12.1	0.04	293 394	73.8	119 160	0.26	2493 2496	34	2500 2495	10 8.6	0	2.117	1.6 0.7	0.1643	0.6	10.69	1.7	0.4722	1.6	0.94
9.1	0.15	146	20.9	62.1	0.15	2583	36	2619	14	1	2.028	1.7	0.1764	0.9	11.99	1.9	0.4929	1.7	0.89
45.1	0.04	282	139	124	0.51	2660	19	2651	11	0	1.958	0.9	0.1798	0.7	12.66	1.1	0.5107	0.9	0.80
29.1	0.08	146	103	65.4	0.31	2698	38	2649	18	-1	1.931	1.7	0.1796	1.1	12.69	2.1	0.5124	1.7	0.33
42.1	0.12	251	168	114	0.69	2741	17	2748	12	0	1.887	0.8	0.1907	0.7	13.93	1.0	0.5298	0.8	0.73
46.1 Granitoid fra	0.34 agments fr	156 rom hemir	120 elagic sec	71.2 liments at	0.80 the North Po	2741 ple site (m	20 easuremer	2/56 hts in thin	18 sections), sample A	1.884 AF07-01 (89°5	0.9 59'10.9'	0.1916 N, 32°19'13.8")	1.1	14.00	1.4	0.5300	0.9	0.63
2_1.1	0.27	214	192	99.8	0.92	nd	nd	2651	21	nd	nd	nd	0.1798	1.2	nd	nd	nd	nd	nd
3 2.1	0.63	692	313	294	0.67	nd	nd	2664	12	nd	nd	nd	0.1834 0.1817	0.7	nd	nd	nd	nd	nd
4_1.1	-0.02	763	407	413	0.55	nd	nd	2943	7.3	nd	nd	nd	0.2149	0.5	nd	nd	nd	nd	nd
4_2.1 4_3.1	2.34	808 2190	414 802	464	0.53	nd nd	nd	2976	10	nd nd	nd nd	nd nd	0.2193 0.1714	0.6	nd	nd nd	nd	nd nd	nd nd
4_4.1	0.20	2411	711	801	0.30	nd	nd	2635	6.0	nd	nd	nd	0.1781	0.4	nd	nd	nd	nd	nd
4_4.2	0.03	3209	984 149	1552	0.32	nd	nd	2633	8.0	nd	nd	nd	0.1778	0.5	nd	nd	nd	nd	nd
4_4.5	0.39	1396	140	661	0.28	nd	nd	2709	10	nd	nd	nd	0.1398	0.6	nd	nd	nd	nd	nd
4_6.1	0.07	1341	223	732	0.17	nd	nd	2954	7.9	nd	nd	nd	0.2163	0.5	nd	nd	nd	nd	nd
4_6.2 4 7.1	0.07	4328	167 294	558 1559	0.16 0.07	nd nd	nd nd	2950 2370	9.2 8.7	nd nd	nd nd	nd nd	0.2159 0.1521	0.6	nd nd	nd nd	nd nd	nd nd	nd nd
4_8.1	0.29	1533	286	608	0.19	nd	nd	2625	9.4	nd	nd	nd	0.1770	0.6	nd	nd	nd	nd	nd
4_8.2	0.06	1723	131	868	0.08	nd	nd	2871	8.1	nd	nd	nd	0.2056	0.5	nd	nd	nd	nd pd	nd
4_9.1	0.09	1486	278	856	0.19	nd	nd	2987	8.2	nd	nd	nd	0.22021	0.5	nd	nd	nd	nd	nd
5_1.1	4.47	2846	1628	1361	0.59	nd	nd	2360	16	nd	nd	nd	0.1512	0.9	nd	nd	nd	nd	nd
5_2.1 5_3.1	2.33	1739 2617	1277 963	867 1271	0.76	nd nd	nd nd	2492 2366	16 12	nd nd	nd nd	nd nd	0.1634	1.0 0.7	nd nd	nd nd	nd nd	nd nd	nd nd
5_4.1	0.07	4111	496	1812	0.12	nd	nd	2221	10	nd	nd	nd	0.1395	0.6	nd	nd	nd	nd	nd

Table 1. Continued.

Analysis	% 206	ppm	ppm	206ma	232 238	(1) 20	⁶ Pb/ ²³⁸ U	(1) ²⁰⁷ P	b/ ²⁰⁶ Pb	%	(1)		(1)		(1)		(1)		err
number Gneissic roo	²⁰⁰ Pb _c	U e Lomono	Th sov Ridae	(Geophys	icists Spur), s	Age (N ample A	1a) ± abs LR07-16 (8)	Age (Ma 3.152° N. 1) ± abs 56.105°F	Dis.	200U/200Pb	±%	²⁰⁷ Pb ²⁰⁰ Pb	±%	207 Pb /200 U	±%	200Pb /200U	±%	corr
65_1.2	2.67	386	356	21.8	0.95	399	4.5	450	380	13	15.63	1.2	0.0560	8.8	0.493	8.9	0.0639	1.2	0.13
65_1.1 65_1.3	3.27	616 579	657 553	35.6	1.10	406	4.3	422	320 400	4	15.37	1.1	0.0552	7.4 9.2	0.495	7.5	0.0650	1.1	0.15
65_2.2	2.47	1572	663	99.2	0.44	446	3.4	532	160	19	13.96	0.8	0.0580	3.9	0.573	3.9	0.0716	0.8	0.20
65_2.3	3.07	1454	622	92.5	0.44	446	3.7	308	230	-31	13.93	0.9	0.0525	5.3	0.519	5.3	0.0717	0.9	0.16
65_2.1 9 2.2	0.20	462	693 325	36.8	0.43	454 571	3.3 4.7	496 570	120 86	9	13.71	0.8	0.0571 0.0591	2.7	0.574	2.8	0.0729	0.8	0.27
9_2.1	0.41	351	211	28.2	0.62	573	5.7	591	200	3	10.75	1.0	0.0597	4.7	0.765	4.8	0.0930	1.0	0.22
9_1.4	0.05	928	49.0	88.9	0.05	681	5.0	645	48	-5	8.972	0.8	0.0612	1.1	0.940	1.4	0.1115	0.8	0.57
9_1.2 9_1.1	0.00	804 704	749	67.7	1.10	684	5.2	682	53	0	8.938	0.8	0.0630	1.0	0.969	1.2	0.1116	0.8	0.62
9_3.2	1.45	126	57.8	12.4	0.47	692	9.2	653	340	-6	8.810	1.4	0.0614	8.0	0.960	8.2	0.1134	1.4	0.17
9_1.3 9_3.1	0.08	313	244 63.2	30.8	0.80	699 708	6.1 7.7	688 679	65 120	-2	8.734 8.616	0.9	0.0624	1.6	0.985	1.8	0.1145	0.9	0.50
4_1.1	0.55	1049	347	175	0.34	1137	7.7	1160	46	2	5.180	0.7	0.0785	1.1	2.088	1.4	0.1929	0.7	0.54
Granitoid fr	agments fr 2 02	om the Po	dvodniko 321	v Basin, sa 416	mple BC-299	9 (81°N, 1 1196	65°E) 7.0	1872	19	57	4 907	0.7	0 1144	1.1	3 217	13	0.2038	0.7	0.52
5.1	0.80	573	77.9	143	0.14	1627	9.9	2085	17	28	3.483	0.7	0.1290	1.0	5.110	1.2	0.2871	0.7	0.58
7.1	1.15	1171 464	372	296	0.33	1643	9.1	2305	17	40	3.446	0.6	0.1464	1.0	5.862	1.2	0.2902	0.6	0.54
mica-5.1	0.95	765	407	277	0.55	2244	15	2615	17	17	2.402	0.8	0.1758	1.0	10.10	1.3	0.4164	0.8	0.60
mica-3.1	0.59	598	195	218	0.34	2267	13	2591	16	14	2.373	0.7	0.1734	1.0	10.08	1.2	0.4214	0.7	0.58
6.2	0.14	972	134	367	0.28	2308	19	2585	10	12	2.323	0.7	0.1728	0.7	10.26	1.4	0.4304	0.7	0.71
7_2.1	0.28	619	285	238	0.48	2377	13	2596	13	9	2.243	0.7	0.1740	0.8	10.69	1.0	0.4458	0.7	0.65
6.1	0.08	577	47.6	232	0.09	2473	15	2637	16	7	2.138	0.7	0.1782	1.0	11.49	1.2	0.4677	0.7	0.60
7_3.1	0.50	415	157	170	0.34	2492	16	2622	13	5	2.1129	0.8	0.1766	0.8	11.53	1.1	0.4090	0.8	0.70
mica-2.1	0.42	341	292	141	0.88	2515	16	2628	14	5	2.096	0.8	0.1774	0.8	11.67	1.2	0.4771	0.8	0.69
1_12.1	1.74	91.0	49.7	7.39	0.56	573	13	644	244	12	10.76	2.4	0.0611	11	0.784	12	0.0929	2.4	0.21
1_15.1	1.47	84.8	72.5	6.99	0.88	583	15	529	318	-9	10.57	2.6	0.0580	14	0.756	15	0.0946	2.6	0.18
1_4.1 1_13.1	0.99	14.0	8.42 140	1.25	1.20	595 604	38 13	-127 513	2004 243	-121	10.34	2.3	0.0437	81	0.583	81	0.0967	2.3	0.08
1_17.1	1.52	79.3	41.7	7.13	0.54	633	16	690	236	9	9.696	2.7	0.0625	11	0.888	11	0.1031	2.7	0.23
1_3.1	0.37	97.2 99.5	60.7 95.4	8.73	0.65	638 643	14 14	826 784	140 215	29	9.605 9.540	2.4	0.0666	6.7 10	0.956	7.1	0.1041	2.4	0.33
1_1.1	0.27	668	20.8	61.6	0.03	655	9.3	589	55	-10	9.345	1.5	0.0596	2.5	0.879	2.9	0.1048	1.5	0.51
1_10.1	3.53	39.0	17.7	4.22	0.47	740	31	581	967	-21	8.220	4.4	0.0594	45	0.996	45	0.1217	4.4	0.10
1_6.1	2.41 0.40	81.8	36.8 70.6	14.8 31.0	0.47	1208	27	953 1780	256	-21	4.850	2.4	0.0709	2.0	2.015	2.8	0.2062	2.4	0.19
1_8.1	0.43	148	158	43.9	1.10	1904	30	1978	36	4	2.911	1.8	0.1215	2.0	5.754	2.7	0.3436	1.8	0.67
1_9.1	0.28	266	219	79.6	0.85	1922	27	2037	24	6	2.879	1.6	0.1256	1.4	6.015	2.1	0.3474	1.6	0.76
1_20.1	0.13	283	72.8	87.5	0.27	1943	27	2038	22	3	2.839	1.6	0.1255	1.2	6.233	2.0	0.3525	1.6	0.49
1_5.1	1.42	98.6	131	32.0	1.37	2041	37	1966	82	-4	2.686	2.1	0.1206	4.6	6.194	5.1	0.3724	2.1	0.41
1_24.1	1.00	522 223	81.6 125	173 74.2	0.16	2087	29 30	2602 2045	21 41	25 -2	2.615	1.7	0.1746	1.3	9.203 6.678	2.1	0.3824	1.7	0.80
1_18.1	0.23	651	122	215	0.19	2096	27	2086	15	0	2.602	1.5	0.1291	0.9	6.840	1.8	0.3843	1.5	0.87
2_5.1	0.63	2444	117	944	0.05	2380	20	2564	8.8	8	2.239	1.0	0.1706	0.5	10.51	1.1	0.4466	1.0	0.88
2_23.2	0.24	897	8.34	352	0.01	2392	20	2563	11	6	2.195	1.0	0.1710	0.6	10.00	1.2	0.4493	1.0	0.85
1_23.2	0.03	1211	15.2	466	0.01	2433	20	2575	13	6	2.181	1.0	0.1717	0.8	10.86	1.3	0.4586	1.0	0.78
1_21.2	0.02	1383	41.0	558 495	0.03	2484 2496	21 22	2565	9.4 12	3	2.127	1.0	0.1707	0.6	11.07	1.2	0.4701	1.0	0.87
1_24.2	0.07	1770	62.8	720	0.04	2496	20	2569	8.7	3	2.115	1.0	0.1711	0.5	11.16	1.1	0.4728	1.0	0.88
2_23.1	0.05	328	133	134	0.42	2506	24	2668	14	6	2.104	1.1	0.1816	0.8	11.90	1.4	0.4752	1.1	0.81
2_11.1	0.04	987	10.2	408	0.03	2532	20	2576	8.8	2	2.101	1.0	0.1733	0.5	11.37	1.1	0.4739	1.0	0.88
2_8.1	0.08	1792	44.4	742	0.03	2533	21	2581	8.1	2	2.078	1.0	0.1724	0.5	11.44	1.1	0.4813	1.0	0.90
1_28.1	0.00	132	413 89.7	55.2 93.5	3.22	2550 2564	43	2696 2670	24 23	6 4	2.061	2.1	0.1848	1.4	12.36	2.5	0.4852	2.1	0.82
1_23.1	0.59	297	156	126	0.54	2574	37	2651	20	3	2.038	1.7	0.1798	1.2	12.16	2.1	0.4907	1.7	0.82
2_9.1	0.29	797	456	338	0.59	2582	22	2676	10	4	2.030	1.0	0.1826	0.6	12.40	1.2	0.4925	1.0	0.87
1_27.1	0.01	540	19.6	230	0.01	2595	35	2592	16	0	2.023	1.5	0.1745	0.9	11.89	1.7	0.4942	1.5	0.73
2_16.2	0.00	285	143	122	0.52	2608	24	2657	12	2	2.005	1.1	0.1805	0.7	12.41	1.3	0.4987	1.1	0.85
1_23.3	0.66	119	38.9	51.5 809	0.34	2609	32	2663	27	-1	2.005	1.5	0.1811 0.1726	1.6	12.45	2.2	0.4988	1.5	0.68
2_16.1	0.64	970	60.6	421	0.06	2626	21	2604	10	-1	1.989	1.0	0.1748	0.6	12.12	1.2	0.5028	1.0	0.85
2_12.1	0.07	1942 325	60.1	840 141	0.03	2627	25 24	2595 2664	15	-1	1.988	1.1	0.1739	0.9	12.06	1.5	0.5031	1.1	0.78
2_19.1 2_4.1	0.04	867	252	377	0.30	2638	24	2663	8.7	1	1.977	1.0	0.1812	0.5	12.63	1.1	0.5052	1.0	0.85
2_17.1	0.14	757	386	330	0.53	2642	21	2673	8.8	1	1.974	1.0	0.1822	0.5	12.73	1.1	0.5066	1.0	0.88
1_21.1	0.63	149 520	69.0 307	65.7 229	0.48	2652	43 22	2682	30 8.8	1	1.965	2.0	0.1832 0.1830	0.5	12.85	2.7	0.5089 0.5108	2.0	0.74
1_11.1	0.09	292	89.5	129	0.32	2673	34	2651	17	-1	1.946	1.5	0.1798	1.0	12.74	1.9	0.5138	1.5	0.83
2_14.2	0.00	127	52.0	56.1	0.42	2676	37	2689	28	0	1.944	1.7	0.1839	1.7	13.05	2.4	0.5145	1.7	0.71
2_18.1 2_20.1	0.40	618	+o.0 369	276	0.55	2694	29	2695	21 10	0	1.941	1.0	0.1840	0.6	13.10	1.8	0.5151	1.0	0.72
2_3.1	0.03	696	168	311	0.25	2697	22	2666	7.8	-1	1.925	1.0	0.1815	0.5	13.00	1.1	0.5196	1.0	0.90
1_20.1 2.24.1	0.71	171 185	88.2 79.2	76.9 83.0	0.53	2699 2706	42 28	2665 2653	27 20	-1 -2	1.923	1.9	0.1814	1.6 1.2	13.01 12.95	2.5 1.8	0.5200	1.9	0.76
1_22.1	0.09	756	278	340	0.38	2712	34	2686	12	-1	1.912	1.6	0.1836	0.7	13.24	1.7	0.5229	1.6	0.91
1_25.1	0.46	195	88.8	88.0	0.47	2716	41	2674	22	-2	1.908	1.9	0.1823	1.3	13.17	2.3	0.5240	1.9	0.82
2_7.1 1_19.1	0.12	384 1329	202	205 605	0.12	2736	23 33	2688	9.5 10	-2	1.891	1.0	0.1839	0.6	13.41	1.2 1.6	0.5288	1.0	0.88
2_22.1	0.06	612	226	279	0.38	2747	25	2688	13	-2	1.882	1.1	0.1839	0.8	13.47	1.4	0.5313	1.1	0.81
2_21.1	0.03	328 1000	127 743	150 459	0.40	2747	24	2691	11 10	-2	1.882	1.1	0.1842	0.7	13.50	1.3	0.5314	1.1	0.85
2_15.1	0.02	1765	615	818	0.36	2781	21	2674	5.8	-4	1.854	0.9	0.1823	0.3	13.56	1.0	0.5393	0.9	0.94
2_2.1	0.03	2173	481	1011	0.23	2791	21	2673	6.1 8.4	-4	1.846	0.9	0.1822	0.4	13.61	1.0	0.5417	0.9	0.93
4_1.4	0.05	743	020	4.0	0.77	2001	43	2019	0.4		1.013	1.0	0.1020	0.5	10.90	1.1	0.5514	1.0	0.09

Table 1. Continued.

Analysis number	% ²⁰⁶ Pbc	ppm U	ppm Th	ppm ²⁰⁶ Pb*	²³² Th/ ²³⁸ U	(1) ²⁰⁶ I Age (Ma	Pb/ ²³⁸ U a) ± abs	(1) ²⁰⁷ Pb Age (Ma)	/ ²⁰⁶ Pb ± abs	% Dis.	(1) ²³⁸ U/ ²⁰⁶ Pb [*]	±%	(1) ²⁰⁷ Pb ^{*/206} Pb [*]	±%	(1) ²⁰⁷ Pb ^{*/235} U	±%	(1) ²⁰⁶ Pb ^{*/238} U	±%	err corr
2_6.1	0.02	1176	678	558	0.60	2836	22	2680	7.8	-5	1.809	1.0	0.1830	0.5	13.95	1.1	0.5527	1.0	0.90
2.1 13.1 4.1 1_1.1 1_2.1 1_1.2 10.1 1.1 1.1 1.1 1.1 3.1 8.1	4.38 3.52 0.60 0.89 4.79 1.61 0.00 0.78 1.56 2.21 0.90	861 2869 949 1450 2834 980 3051 687 1818 492 703	402 353 240 305 856 386 222 457 170 174 302	184 714 228 356 795 272 876 205 582 165 241	0.48 0.13 0.26 0.22 0.31 0.41 0.08 0.69 0.10 0.37 0.44	1353 1576 1580 1607 1731 1774 1859 1909 2007 2071 2145	16 7.7 9.0 9.0 20 21 8.8 9.6 9.7 59 15	2604 2233 2603 2246 2512 2267 2270 2617 2289 2636 2609	20 26 9.5 17 82 17 5.1 10 13 24 9.5	92 42 65 40 45 28 22 37 14 27 22	4.214 3.575 3.593 3.534 3.247 3.157 2.992 2.894 2.725 2.618 2.524	1.3 0.6 0.7 0.6 1.3 1.3 0.5 0.6 0.6 3.3 0.8	0.1748 0.1405 0.1746 0.1415 0.1649 0.1432 0.1435 0.1761 0.1451 0.1782 0.1753	$ \begin{array}{c} 1.2\\ 1.5\\ 0.6\\ 1.0\\ 4.9\\ 1.0\\ 0.3\\ 0.6\\ 0.8\\ 1.4\\ 0.6\\ \end{array} $	5.630 5.362 6.687 5.524 7.030 6.260 6.613 8.367 7.307 9.310 9.546	1.8 1.6 0.9 1.2 5.0 1.7 0.6 0.8 0.9 3.6 1.0	0.2335 0.2769 0.2777 0.2830 0.3080 0.3168 0.3342 0.3446 0.3653 0.3790 0.3949	$\begin{array}{c} 1.3 \\ 0.6 \\ 0.7 \\ 0.6 \\ 1.3 \\ 1.3 \\ 0.5 \\ 0.6 \\ 0.6 \\ 3.3 \\ 0.8 \\ 0.8 \end{array}$	0.73 0.34 0.75 0.55 0.26 0.81 0.88 0.69 0.60 0.92 0.83
5.1 7.1 6.1 9.1 12.1 Grapitoid fr	1.86 0.29 1.33 0.00 0.26	499 565 303 479 1042	472 105 115 526	174 209 121 196 442	0.82 0.86 0.36 0.25 0.52	2139 2301 2423 2512 2578 79°N 178°	10 12 14 13 13	2627 2630 2624 2613 2610	19 7.9 14 7.3 14	22 14 8 4 1	2.499 2.328 2.181 2.098 2.032	0.9 0.6 0.7 0.6 0.6	0.1772 0.1775 0.1770 0.1758 0.1754	0.5 0.9 0.4 0.9	9.720 10.50 11.13 11.55 11.89	1.4 0.8 1.1 0.8 1.0	0.3978 0.4290 0.4563 0.4766 0.4918	0.9 0.6 0.7 0.6 0.6	0.81 0.79 0.62 0.82 0.59
3.1 1.1 7.1 5.1 10.1 4.1 12.1 8.1 2.2 6.1 9.1 11.1 Granitoid fra	0.00 0.64 1.52 2.26 0.46 0.00 0.76 1.93 0.13 0.00 0.00 0.00 0.00 0.17 agment fro	466 716 389 1147 754 1107 577 561 428 423 533 352 1079 200 the Mee	375 435 405 349 606 832 304 206 545 402 161 296 1451 ndeleev R	96.6 151 96.1 328 222 340 192 192 192 149 150 198 147 497 ise, sampl	0.83 0.63 1.07 0.31 0.83 0.54 0.54 0.38 1.32 0.98 0.31 0.87 1.39 e BC-260 (80	1393 1407 1607 1817 1890 1971 2098 2123 2194 2228 2313 2557 2762 °N, 179°30	13 12 16 15 16 16 16 18 18 19 20 19 22 23	2599 2325 2331 2330 2339 2602 2321 2353 2615 2625 2624 2622 2323	22 18 28 13 8.2 27 24 11 10 9.4 11 9.4	86 65 45 28 24 32 11 11 19 18 13 3 -16	4.144 4.099 3.533 3.072 2.936 2.795 2.600 2.564 2.467 2.422 2.317 2.054 1.869	$ \begin{array}{c} 1.0\\ 1.0\\ 1.1\\ 1.0\\ 1.0\\ 1.0\\ 1.0\\ 1.0\\$	$\begin{array}{c} 0.1742\\ 0.1482\\ 0.1487\\ 0.1486\\ 0.1494\\ 0.1746\\ 0.1479\\ 0.1506\\ 0.1760\\ 0.1760\\ 0.1770\\ 0.1769\\ 0.1767\\ 0.1480\\ \end{array}$	$\begin{array}{c} 1.3 \\ 1.0 \\ 1.6 \\ 1.0 \\ 0.7 \\ 0.5 \\ 1.6 \\ 1.4 \\ 0.7 \\ 0.6 \\ 0.6 \\ 0.7 \\ 0.6 \end{array}$	5.796 4.984 5.800 6.669 7.017 8.610 7.840 8.100 9.840 10.08 10.53 11.86 10.92	1.6 1.4 2.0 1.4 1.2 1.1 1.8 1.7 1.2 1.2 1.1 1.2 1.2 1.2	0.2413 0.2440 0.2831 0.3255 0.3406 0.3577 0.3846 0.3901 0.4054 0.4129 0.4316 0.4870 0.5350	$\begin{array}{c} 1.0\\ 1.0\\ 1.1\\ 1.0\\ 1.0\\ 1.0\\ 1.0\\ 1.0\\$	0.61 0.68 0.58 0.68 0.80 0.89 0.54 0.57 0.83 0.86 0.87 0.84 0.88
4.1r 4.1 7_3.1 5.1 4_2.1 3.1 6.1 1.2 7_2.1 1.1 7.1	2.40 2.66 3.12 1.33 1.30 1.21 2.35 1.86 1.16 0.21 1.42	12073 9257 3431 5145 4308 4915 3530 2573 3043 2077 1392	4824 4296 4078 5481 1904 686 1132 273 6097 49.7 270	189 224 144 226 236 280 230 289 348 359 245	0.41 0.48 1.23 1.10 0.46 0.14 0.33 0.11 2.07 0.02 0.20	114 174 299 317 393 410 461 777 796 1178 1183	6.4 1.6 6.3 2.6 4.7 3.3 6.3 21 7.9 7.1 18	1744 1822 1705 1805 1850 1922 1892 1931 1968 1933 1971	38 47 71 51 24 23 97 37 30 30 30 35	1434 947 471 469 371 369 311 148 147 64 67	56.20 36.55 21.08 19.82 15.91 15.24 13.48 7.790 7.599 4.985 4.957	5.7 0.9 2.1 0.9 1.2 0.8 1.4 2.9 1.0 0.7 1.7	0.1067 0.1113 0.1044 0.1103 0.1131 0.1177 0.1157 0.1183 0.1208 0.1184 0.1210	2.1 2.6 3.8 2.8 1.3 5.4 2.1 1.7 1.7 2.0	0.262 0.420 0.683 0.767 0.980 1.065 1.183 2.091 2.190 3.275 3.361	6.1 2.8 4.4 2.9 1.8 1.5 5.6 3.6 2.0 1.8 2.6	0.0178 0.0274 0.0474 0.0504 0.0628 0.0656 0.0741 0.1282 0.1315 0.2006 0.2014	$5.7 \\ 0.9 \\ 2.1 \\ 0.9 \\ 1.2 \\ 0.8 \\ 1.4 \\ 2.9 \\ 1.0 \\ 0.7 \\ 1.7 $	0.94 0.33 0.49 0.29 0.69 0.54 0.25 0.81 0.53 0.36 0.64
$\begin{array}{c} \text{metagabbt}\\ 31_1.1\\ 31_1.2\\ 31_2.1\\ 31_3.1\\ 31_3.2\\ 31_4.1\\ 31_5.1\\ 31_6.1\\ 31_6.2\\ 31_7.1\\ 32_1.1 \end{array}$	0.70 0.00 9.57 16.3 10.4 7.73 0.00 0.13 1.33 0.15 0.09	168 253 412 486 442 1485 95.2 709 1905 440 1895	67.1 145 841 1152 748 1476 16.0 224 102 130 2294	17.9 28.8 83.2 107 149 370 28.7 317 815 197 193	0.41 0.59 2.11 2.45 1.75 1.03 0.17 0.33 0.06 0.31 1.25	750 801 1242 1248 1946 1528 1940 2700 2577 2701 nd	9.4 8.7 14 16 27 15 24 22 20 22 nd	818 811 1904 1989 1958 1923 1962 2661 2652 2673 776	102 37 85 98 94 43 30 8 9 10 20	9 1 53 59 1 26 1 -1 3 -1 nd	8.101 7.561 4.705 4.682 2.838 3.738 2.848 1.923 2.034 1.921 nd	1.3 1.2 1.3 1.4 1.6 1.1 1.4 1.0 1.0 1.0 nd	0.0664 0.0661 0.1166 0.1222 0.1201 0.1178 0.1204 0.1809 0.1799 0.1822 0.0651	4.9 1.8 4.7 5.5 5.3 2.4 1.7 0.5 0.5 0.6 0.9	1.130 1.206 3.415 3.599 5.836 4.344 5.828 12.97 12.19 13.08 nd	5.1 2.1 4.9 5.7 5.5 2.6 2.2 1.1 1.1 1.2 nd	0.1234 0.1323 0.2125 0.2136 0.3524 0.2675 0.3511 0.5201 0.4915 0.5205 nd	1.3 1.2 1.3 1.4 1.6 1.1 1.4 1.0 1.0 1.0 nd	0.26 0.54 0.26 0.24 0.29 0.43 0.65 0.90 0.88 0.86 nd
Metagabbro 1.1 2.1 2.2 3.1 3.2 4.1 5.1 6.1 7.1 7.2 8.1 9.1 10.1 11.1	dolerite 1 0.46 0.55 0.48 0.43 0.67 1.14 3.52 3.96 0.86 0.89 1.80 11.1 13.4	fragments 6991 1067 2090 1499 1063 1006 12374 1729 2938 2884 1559 3550 934 964	from the l 24825 557 2055 1076 611 263 36736 1117 3131 4809 937 8734 1467 1532	Mendeleev 445 99.8 207 147 105 109 806 198 240 221 173 220 132 117 221 220 221 220 221 220 221 220 220	v Rise (measu 3.67 0.54 1.02 0.74 0.59 0.27 3.07 0.67 1.10 1.72 0.62 2.54 1.62 1.64	urements i nd nd nd nd nd nd nd nd nd nd nd nd nd	n thin sect nd nd nd nd nd nd nd nd nd nd nd nd nd	tions), sam 720 828 786 800 816 804 717 836 733 723 859 760 839 397	ple AF05 15 30 21 28 56 33 48 268 123 29 60 39 210 610	5-26 (79°N nd nd nd nd nd nd nd nd nd nd nd nd nd	, 178°W) nd nd nd nd nd nd nd nd nd nd nd nd nd	nd nd nd nd nd nd nd nd nd nd nd nd nd	0.0633 0.0667 0.0654 0.0658 0.0663 0.0659 0.0633 0.0669 0.0633 0.0669 0.0634 0.0670 0.0646 0.0670 0.0546	0.7 1.4 1.0 1.3 2.7 1.6 2.3 13 5.8 1.4 2.9 1.8 10 27	nd nd nd nd nd nd nd nd nd nd nd nd nd n	nd nd nd nd nd nd nd nd nd nd nd nd nd	nd nd nd nd nd nd nd nd nd nd nd nd nd n	nd nd nd nd nd nd nd nd nd nd nd nd nd	nd nd nd nd nd nd nd nd nd nd nd nd nd n
spandstone f 2_6.1 1_8.1 1_9.1 1_10.1 2_6.2 1_41.1 2_17.1 2_25.1 2_16.1 1_1.1 2_23.1 2_12.1 2_9.1 2_28.1 2_30.1 2_22.1 2_11.1	aragments 1.20 0.32 0.01 0.34 0.38 0.42 0.23 0.65 0.46 0.22 0.11 0.20 0.011 0.25 0.05 0.19 0.03	rrom the N 3429 887 252 833 264 2056 110 421 16.6 419 126 419 126 419 126 235 709 323 370 121 385	nendeleev 1583 496 119 388 1036 72.1 1036 93.6 87.9 126 42.0 57.8 97.2 80.5 270 80.5 270 87.3 54.8 172 244	Hise, sam 104 29.4 9.86 32.5 10.6 88.2 5.02 24.5 0.97 25.8 7.98 15.7 9.21 81.7 9.21 81.7 9.21 81.7 9.21 81.7 9.21 81.7 9.21 81.7 9.21 81.7 9.21 81.7 9.21 81.7 9.21 81.7 9.21 81.7 9.20.7 32.9 101 47.4 20.3 64.6 58.6	pres Ar00-05 0.48 0.58 0.49 0.49 0.49 0.74 0.52 0.68 0.39 0.01 0.23 0.72 0.76 0.49 0.10 0.65 0.35 0.35 0.39 0.31 0.22 0.47 0.47 0.47	 (1, 2) (82' 221 243 284 286 291 313 327 421 422 444 456 648 734 917 971 9916 1069 1145 1150 1162 	10 N, 178 0.8 1.8 4.6 2.1 1.0 5.3 3.0 11 3.0 4.3 7.0 7.9 3.7 6.7 5.1 4.3 5.2 5.1 9.3 5.0 4.6	w) 1561 119 -82 201 -8 1675 -451 446 178 450 233 770 888 1032 912 888 1032 1072 1182 1157	$\begin{array}{c} 24\\ 53\\ 206\\ 71\\ 193\\ 13\\ 611\\ 66\\ 478\\ 86\\ 478\\ 86\\ 74\\ 119\\ 73\\ 33\\ 35\\ 36\\ 21\\ 25\\ 24\\ 52\\ 15\\ 15\\ 15\\ \end{array}$	$\begin{array}{c} 606\\ -51\\ -129\\ -30\\ -103\\ 435\\ -238\\ 6\\ -58\\ 6\\ -58\\ -1\\ -49\\ -19\\ 5\\ -4\\ -2\\ -9\\ 4\\ 2\\ 0\\ 3\\ 1\\ 0\end{array}$	28.68 25.91 22.04 21.99 21.47 20.10 18.89 14.82 14.78 14.04 13.63 9.404 8.291 6.540 6.464 6.150 6.018 5.857 5.544 5.143 5.143	$\begin{array}{c} 0.3\\ 0.8\\ 1.7\\ 0.7\\ 1.1\\ 0.3\\ 1.6\\ 0.7\\ 2.7\\ 0.7\\ 1.0\\ 1.1\\ 1.1\\ 1.1\\ 0.4\\ 0.8\\ 0.6\\ 0.5\\ 0.6\\ 0.5\\ 0.6\\ 0.5\\ 0.4\\ \end{array}$	0.0967 0.0515 0.0521 0.0521 0.1028 0.0520 0.0559 0.0509 0.0649 0.0684 0.0684 0.0684 0.0684 0.0684 0.0737 0.0731 0.0751 0.0751 0.0794 0.0784	$\begin{array}{c} 1.3\\ 1.9\\ 3.7\\ 1.8\\ 3.7\\ 0.7\\ 5.7\\ 3.0\\ 20\\ 3.9\\ 3.2\\ 3.4\\ 3.5\\ 1.6\\ 1.7\\ 1.8\\ 1.0\\ 1.3\\ 1.2\\ 2.6\\ 0.7\\ 0.8 \end{array}$	0.465 0.270 0.322 0.335 0.705 0.520 0.463 0.549 0.549 0.549 0.549 0.514 0.912 1.079 1.443 1.481 1.539 1.688 1.739 1.868 2.128 2.111 2.135	1.3 2.0 4.1 1.9 3.8 0.8 6.0 3.1 21 3.9 3.4 3.6 3.7 1.7 1.8 1.8 1.1 1.4 1.3 2.8 0.9 0.9	0.0349 0.0386 0.0454 0.0455 0.0466 0.0498 0.0529 0.0675 0.0677 0.0712 0.07712 0.07712 0.07712 0.0703 0.1206 0.1529 0.1547 0.1662 0.1662 0.1662 0.1707 0.1804 0.1944 0.1944	$\begin{array}{c} 0.3\\ 0.8\\ 1.7\\ 0.7\\ 1.1\\ 0.3\\ 1.6\\ 0.7\\ 2.7\\ 0.7\\ 1.0\\ 1.1\\ 1.1\\ 0.4\\ 0.8\\ 0.6\\ 0.5\\ 0.6\\ 0.5\\ 0.9\\ 0.5\\ 0.4\\ \end{array}$	$\begin{array}{c} 0.26\\ 0.37\\ 0.40\\ 0.38\\ 0.29\\ 0.42\\ 0.28\\ 0.24\\ 0.13\\ 0.18\\ 0.29\\ 0.32\\ 0.31\\ 0.42\\ 0.31\\ 0.40\\ 0.40\\ 0.40\\ 0.32\\ 0.54\\ 0.49\\ \end{array}$

Table 1. Continued.

Analysis	%	ppm	ppm	ppm		(1) 206	Pb/ ²³⁸ U	(1) ²⁰⁷ Ph	/ ²⁰⁶ Pb	%	(1)		(1)		(1)		(1)		err
number	²⁰⁶ Pb _c	U	Th	²⁰⁶ Pb*	$^{232}{\rm Th}/^{238}{\rm U}$	Age (Ma	a) ± abs	Age (Ma)	± abs	Dis.	$^{238}\text{U}/^{206}\text{Pb}^{*}$	±%	$^{207}{\rm Pb}^{*/206}{\rm Pb}^{*}$	±%	207Pb*/235U	±%	$^{206}{Pb}^{*/238}U$	±%	corr
2_10.1	0.03	155	55.6	28.3	0.37	1239	7.5	1135	23	-8	4.718	0.7	0.0775	1.1	2.266	1.3	0.2120	0.7	0.50
1_6.1	0.12	310	80.9	57.9	0.27	1263	9.1	1707	26	35	4.612	0.8	0.1057	1.3	3.160	1.5	0.2168	0.8	0.53
2_26.1	0.08	164	70.0	33.7	0.44	1381	10	1387	30	0	4.185	0.8	0.0882	1.6	2.906	1.8	0.2390	0.8	0.47
2_14.1	0.06	57.0	30.0	45.0	0.44	1594	/.8	1500	43	-2	4.142	1.2	0.0870	2.3	2.890	2.6	0.2414	1.2	0.57
1 3.1	0.19	372	510	96.2	1.41	1691	11	1839	17	9	3.330	0.7	0.1130	0.8	4.681	1.1	0.3003	0.7	0.66
2_3.1	0.02	365	224	95.6	0.63	1714	9.4	1849	13	8	3.282	0.6	0.1131	0.7	4.749	1.0	0.3047	0.6	0.66
2_13.1	0.05	92.7	44.2	24.7	0.49	1741	14	1744	19	0	3.226	0.9	0.1067	1.0	4.562	1.4	0.3100	0.9	0.65
2_29.1	0.17	107	41.8	29.3	0.40	1778	13	1918	27	8	3.148	0.8	0.1174	1.5	5.144	1.7	0.3177	0.8	0.48
1_5.1		27.7	15.0	7.68	0.56	1795	36	1648	81	-8	3.080	2.3	0.1100	3.9	4.924	4.6	0.3247	2.3	0.50
2_24.1	0.02	215	35.0	37.9	0.27	1808	12	1822	22	1	3.089	0.7	0.1114	1.2	4.970	1.4	0.3237	0.7	0.52
2_2.1	0.03	78.5	62.8	23.3	0.83	1912	17	1842	27	-4	2.896	1.0	0.11228	1.5	5.361	1.8	0.3453	1.0	0.57
2_20.1		335	150	101	0.46	1938	9.0	1933	13	0	2.851	0.5	0.1184	0.8	5.727	0.9	0.3508	0.5	0.58
2_5.1	0.01	255	162	77.4	0.66	1951	11	1997	11	2	2.830	0.7	0.1228	0.6	5.982	0.9	0.3534	0.7	0.72
1_7.1	0.19	207	136	63.8	0.68	1975	15	1979	19	0	2.785	0.9	0.1225	1.0	6.067	1.3	0.3591	0.9	0.67
2_8.1	0.01	115	76.0	35.5	0.68	1976	12	1991	14	1	2.788	0.7	0.1224	0.8	6.051	1.0	0.3587	0.7	0.67
2_17.2	0.14	151	55.6 78.1	23.5	0.76	1989	15	1993	19	0	2.767	0.9	0.1225	1.1	6.103	1.4	0.3613	0.9	0.63
2_27.1	0.00	527	345	226	0.55	2514	8.0	2500	44	4	2.097	0.7	0.1703	0.8	12.87	0.5	0.4770	0.7	0.82
Sandstone f	ragments	from the M	Aendeleev	Rise, sam	ple AF00-10	(82°04'N,	179°59'W)							0.10		0.10			
15.2	1.71	1929	980	64.2	0.53	241	1.0	552	91	129	26.24	0.4	0.0586	4.2	0.308	4.2	0.0381	0.4	0.10
15.1	1.60	1459	693	60.6	0.49	300	1.4	567	68	89	21.02	0.5	0.0590	3.1	0.387	3.1	0.0476	0.5	0.15
26.1	1.00	1008	172	49.0	0.18	352	1.7	393	74	12	17.83	0.5	0.0545	3.3	0.421	3.4	0.0561	0.5	0.15
5.1	0.66	900	963	47.9	1.11	385	2.5	350	70	-9	16.24	0.7	0.0548	2.6	0.454	3.2	0.0615	0.7	0.21
3.1	1.44	436	209	34.3	0.48	490 564	4.9	481	39	-15	10.91	0.9	0.0923	1.8	0.742	2.7	0.0790	0.9	0.50
20.1	1.15	1799	1466	205	0.84	795	4.2	1539	63	94	7.619	0.6	0.0955	3.3	1.729	3.4	0.1313	0.6	0.17
1.1	0.07	774	373	106	0.50	949	5.5	920	25	-3	6.297	0.6	0.0704	1.0	1.525	1.4	0.1587	0.6	0.46
22.1	0.64	826	32.8	115	0.04	965	3.9	1186	27	23	6.193	0.4	0.0796	1.4	1.772	1.5	0.1615	0.4	0.30
25.1	0.23	151	112	21.0	0.77	965	7.4	995	54	3	6.193	0.8	0.0723	2.7	1.610	2.8	0.1615	0.8	0.29
27.1	0.01	316	159	52.1	0.52	1132	7.0	1012	25	2	5.208	0.7	0.0783	1.2	2.073	1.4	0.1920	0.7	0.48
21.1	1.24	42.6	73.5	10.2	1.78	1612	19	1668	75	-12	3.520	1.3	0.1024	4.1	4.010	4.3	0.2841	1.3	0.31
29.1	0.01	293	132	73.4	0.47	1651	7.8	1664	16	1	3.427	0.5	0.1022	0.9	4.110	1.0	0.2918	0.5	0.52
11.1	0.28	44.8	52.5	11.9	1.21	1731	25	1738	46	0	3.246	1.6	0.1064	2.5	4.518	3.0	0.3080	1.6	0.54
30.1		66.7	35.7	17.8	0.55	1739	16	1748	32	1	3.229	1.0	0.1069	1.7	4.565	2.0	0.3097	1.0	0.51
28.1	0.05	132	44.3	36.1	0.35	1785	12	1821	22	2	3.135	0.8	0.1113	1.2	4.896	1.4	0.3190	0.8	0.52
16.1	0.10	436	113	121	0.27	1801	6.3	1858	12	3	3.103	0.4	0.1136	0.7	5.048	0.8	0.3223	0.4	0.53
4.1	0.10	272	07.7 76.0	78.9	0.45	1854	23 14	1932	24	4	2.980	0.9	0.1184	1.4	5 322	2.0	0.3349	1.4	0.75
10.1		193	86.9	55.5	0.47	1866	15	1841	20	-1	2.976	0.9	0.1137	1.3	5.209	1.4	0.3357	0.9	0.64
24.1	0.09	634	317	183	0.52	1868	5.8	1966	17	5	2.975	0.4	0.1207	0.9	5.593	1.0	0.3361	0.4	0.36
9.2		98.9	32.9	29.2	0.34	1900	20	1877	27	-1	2.910	1.2	0.1169	1.4	5.424	1.9	0.3427	1.2	0.62
13.1	0.27	133	59.0	39.6	0.46	1914	15	1982	23	4	2.893	0.9	0.1218	1.3	5.804	1.6	0.3456	0.9	0.58
14.1	0.09	82.8	58.4	24.8	0.73	1927	16	1962	25	2	2.870	1.0	0.1204	1.4	5.784	1.7	0.3484	1.0	0.57
19.1	0.15	376	348	95.0	0.47	1957	86	1937	12	2	2.810	0.8	0.1207	0.8	5.982	0.9	0.3550	0.8	0.07
12.1	0.35	207	110	63.4	0.55	1962	11	1980	23	1	2.812	0.7	0.1216	1.3	5.962	1.4	0.3557	0.7	0.46
9.1		412	59.5	128	0.15	1988	12	1956	12	-2	2.768	0.7	0.1205	0.7	5.974	1.0	0.3612	0.7	0.72
18.1	0.05	381	183	120	0.50	2018	7.3	2014	11	0	2.720	0.4	0.1239	0.6	6.282	0.7	0.3676	0.4	0.58
23.1	0.19	52.9	22.9	17.0	0.45	2042	19	2034	34	0	2.684	1.1	0.1254	1.9	6.440	2.2	0.3726	1.1	0.50
0.1 Sandstone f	ragments	575 from the M	103 Aendeleev	Rise sam	0.45 ple AF05-11	(78°55'N.	177°40'W)	2706	8.0	-1	1.905	1.2	0.1862	0.5	15.47	1.5	0.3234	1.2	0.93
4_20.1	0.00	22.2	19.2	0.83	0.89	274	20	352	446	28	23.01	7.6	0.0535	20	0.321	21	0.0435	7.6	0.360
4_20.2	0.00	27.6	22.5	1.10	0.84	292	16	492	316	69	21.61	5.5	0.0570	14	0.364	15	0.0463	5.5	0.357
6_3.1	0.63	1984	3426	150	1.78	541	3.4	1598	33	195	11.43	0.7	0.0986	1.7	1.189	1.9	0.0875	0.7	0.353
6_5.1	0.74	2091	926	169	0.46	577	3.0	1489	27	158	10.68	0.5	0.0931	1.4	1.201	1.5	0.0936	0.5	0.360
6.9.1	0.69	1056	945 945	87.3	0.29	588	3.4	1360	41 52	95	10.00	0.0	0.0771	2.1	0.997	2.1	0.0956	0.0	0.286
6 8.1	0.65	1056	386	96.3	0.38	646	4.1	1100	52	70	9.483	0.7	0.0762	2.6	1.108	2.7	0.1055	0.7	0.249
6_1.1	0.23	1348	968	124	0.74	655	3.3	1224	26	87	9.344	0.5	0.0811	1.3	1.197	1.4	0.1070	0.5	0.364
6_7.1	0.90	478	322	46.0	0.70	679	6.5	1388	78	104	9.002	1.0	0.0883	4.0	1.352	4.2	0.1111	1.0	0.242
6_20.1	0.23	870	178	84.5	0.21	688	4.6	1192	37	73	8.874	0.7	0.0798	1.9	1.240	2.0	0.1127	0.7	0.351
4 2 1	1.95	106	35.9	16.2	0.80	1037	21	1032	216	0	5 731	2.2	0.0748	11	1.091	11	0.1040	2.2	0.219
6_21.1	0.28	670	248	102	0.38	1048	6.6	1067	37	2	5.665	0.7	0.0750	1.9	1.824	2.0	0.1765	0.7	0.344
6_18.1	1.00	167	61.1	25.9	0.38	1057	13	1053	139	0	5.611	1.3	0.0744	6.9	1.828	7.0	0.1782	1.3	0.190
6_4.1	1.88	77.7	45.5	12.2	0.61	1063	18	1077	184	1	5.578	1.8	0.0753	9.2	1.861	9.3	0.1793	1.8	0.197
4_17.1	3.91	29.4	22.5	4.89	0.79	1100	32	1242	331	13	5.377	3.2	0.0818	17	2.098	17	0.1860	3.2	0.185
4_21.1	0.41	140	09.4 83.4	25.5	0.49	1105	14	1121	52	6	5 346	1.1	0.0770	2.6	2.035	29	0.1871	1.1	0.316
6 14.1	0.02	159	93.0	25.8	0.60	1114	12	1122	75	1	5.301	1.2	0.0770	3.8	2.004	3.9	0.1886	1.4	0.305
6_25.1	0.75	157	89.1	25.9	0.59	1122	14	1208	103	8	5.260	1.3	0.0804	5.2	2.108	5.4	0.1901	1.3	0.248
4_8.1	3.02	21.7	9.68	3.70	0.46	1138	38	1546	257	36	5.181	3.7	0.0959	14	2.552	14	0.1930	3.7	0.260
6_6.1	0.57	164	118	28.5	0.74	1177	13	1264	87	7	4.991	1.2	0.0828	4.5	2.286	4.6	0.2004	1.2	0.260
6_2.1	0.61	427	133	74.8	0.32	1190	8.7	1220	41	3	4.931	0.8	0.0810	2.1	2.264	2.2	0.2028	0.8	0.357
4_14.1	0.17	185	63.9	34.7	0.36	1229	13	1333	40	5	4.701	1.1	0.0823	2.1	2.588	2.5	0.2101	1.1	0.411
4_10.1	5.62	25.7	27.7	5.17	1.11	1288	48	1375	499	7	4.522	4.1	0.0876	26	2.672	26	0.2211	4.1	0.158
6_13.1	0.00	130	52.4	25.1	0.42	1302	15	1356	45	4	4.469	1.2	0.0868	2.3	2.677	2.6	0.2237	1.2	0.473
6_16.1	2.16	29.9	12.8	5.92	0.44	1311	36	1380	298	5	4.436	3.1	0.0879	16	2.732	16	0.2254	3.1	0.193
6_11.1	0.66	31.2	18.3	6.23	0.60	1337	30	1466	136	10	4.338	2.5	0.0919	7.2	2.922	7.6	0.2305	2.5	0.325
4_5.1	1.32	48.2	18.7	9.68	0.49	1387	55 34	1397	120	15	4.107	2.6	0.0985	8.4 12	3.201	8.8 12	0.2400	2.6	0.299
6 23.1	0.96	95.8	42.3	20.3	0.46	1411	21	1414	93	0	4.086	1.7	0.0895	4.9	3.019	5.1	0.2447	1.7	0.325
4_1.1	0.92	39.9	33.0	8.52	0.85	1418	32	1580	115	11	4.063	2.5	0.0976	6.2	3.313	6.7	0.2461	2.5	0.378
4_15.1	0.41	178	77.5	37.7	0.45	1419	14	1422	48	0	4.060	1.1	0.0898	2.5	3.051	2.7	0.2463	1.1	0.390
4_16.1	0.00	134	48.6	28.5	0.38	1427	16	1463	44	3	4.037	1.2	0.0918	2.3	3.135	2.6	0.2477	1.2	0.471
4_7.1	0.38	92.9	55.5 87 2	20.0	0.62	1436	20	1448	60 39	1	4.007	1.6	0.0911	3.2	3.133	3.5	0.2495	1.6	0.446
+_13.1 6 12 1	0.14	138	69.2	30.1	0.40	1455	14	1441	48	0	3.948	1.1	0.0907	2.0	3.120	2.3 2.8	0.2499	1.1	0.439
4 9.1	1.04	112	94.8	24.6	0.88	1461	19	1609	73	10	3.932	1.4	0.0992	3.9	3.479	4.2	0.2543	1.4	0.344
4_12.1	0.63	93.1	101	20.9	1.12	1489	21	1599	78	7	3.848	1.6	0.0986	4.2	3.535	4.5	0.2599	1.6	0.352

| Analysis
number | %
²⁰⁶ Pb | ppm
U | ppm
Th | ppm
²⁰⁶ Pb*

 | ²³² Th/ ²³⁸ U | (1) ²⁰⁶ | ?b / ²³⁸ U | (1) ²⁰⁷ Pb | / ²⁰⁶ Pb
+ abs | %
Dis. | (1)
²³⁸ U/ ²⁰⁶ Pb [*] | +%
 | (1)
²⁰⁷ Pb ^{*/206} Pb [*] | +% | (1)
²⁰⁷ Pb ^{*/235} U
 | +% | (1)
²⁰⁶ Pb ^{*/238} U | +%
 | err |
|--|--|--|--
--
---|---|---|--|--|---|---
---|--|--
--|---
--|--|---
--|
| 4_11.1 | 1.10 | 19.7 | 7.62 | 4.45

 | 0.40 | 1489 | 43 | 1839 | 161 | 24 | 3.848 | 3.3
 | 0.1124 | 8.9
 | 4.030 | 9.5 | 0.2599 | 3.3
 | 0.344 |
| 4_5.1
4_19.1 | 0.10 | 86.0
250 | 51.4
221 | 19.7
61.9

 | 0.62 | 1509
1632 | 13 | 1594
1644 | 30 | 6 | 3.791
3.471 | 1.7
0.9
 | 0.0984 0.1011 | 5.8
1.6
 | 4.016 | 6.0
1.8 | 0.2638 0.2881 | 0.9
 | 0.286
0.478 |
| 6_10.1 | 0.26 | 146 | 244 | 37.3

 | 1.72 | 1668 | 19 | 1730 | 78 | 4 | 3.386 | 1.3
 | 0.1059 | 4.3
 | 4.312 | 4.4 | 0.2954 | 1.3
 | 0.285 |
| 6_26.1
6_22.1 | 0.11 | 201 | 156 | 86.2

 | 0.37 | 2605 | 22 | 2624 | 24
16 | -2 | 2.926 | 0.7
 | 0.1130 | 1.3
0.9
 | 5.328 | 1.5
1.4 | 0.3418 0.4980 | 1.0
 | 0.460 |
| 4_18.1 | 0.26 | 213 | 155 | 95.4

 | 0.75 | 2697 | 19 | 2707 | 17 | 0 | 1.925 | 0.8
 | 0.1860 | 1.0
 | 13.32 | 1.3 | 0.5195 | 0.8
 | 0.630 |
| Sandstone f | ragments | from the N | J /.4
Aendeleev | Rise, sam

 | ole AF05-14 | (79°N, 172 | 27
°W) | 2724 | 23 | -1 | 1.878 | 1.2
 | 0.1879 | 1.5
 | 13.79 | 1.9 | 0.3324 | 1.2
 | 0.019 |
| 4_4.1
4_2.1 | 7.30
0.09 | 678
723 | 166
314 | 40.3
40.6

 | 0.25
0.45 | 400
408 | 4.1
2.8 | 548
409 | 210
31 | 37
0 | 15.60
15.30 | 1.1
0.7
 | 0.0585 0.0549 | 9.5
1.4
 | 0.516
0.495 | 9.5
1.5 | 0.0639 0.0653 | 1.1
0.7
 | 0.11
0.46 |
| 2_23.1 | 0.00 | 357 | 338 | 22.1

 | 0.98 | 449 | 4.2 | 389 | 44 | -13 | 13.88 | 1.0
 | 0.0544 | 2.0
 | 0.541 | 2.2 | 0.0721 | 1.0
 | 0.44 |
| 2_9.1
4_3.1 | 0.06 | 578 | 456
182 | 49.2
39.0

 | 0.60 | 457
486 | 3.6 | 432
519 | 37
47 | -6
7 | 13.61 | 0.8
 | 0.0555 | 2.2
 | 0.562 | 2.3 | 0.0735 | 0.8
 | 0.44 |
| 4_6.1 | 0.08 | 517 | 243 | 45.7

 | 0.49 | 630 | 4.4 | 620 | 26 | -2 | 9.734 | 0.7
 | 0.0605 | 1.2
 | 0.856 | 1.4 | 0.1027 | 0.7
 | 0.51 |
| 2_3.1
2_26.1 | 0.00 | 1372 | 1700 | 132

 | 1.28 | 682 | 6.0 | 1175 | 48
20 | -4
72 | 9.320
8.960 | 0.9
 | 0.0603 | 1.0
 | 1.217 | 1.4 | 0.1051 | 0.9
 | 0.46 |
| 2_12.1 | 2.37 | 707 | 71.0 | 81.4

 | 0.10 | 791 | 7.4 | 1423 | 55
50 | 80 | 7.649 | 1.0
 | 0.0898 | 2.9
 | 1.617 | 3.0 | 0.1305 | 1.0
 | 0.33 |
| 2_21.1 | 0.05 | 567 | 336 | 87.2

 | 0.61 | 1061 | 7.9 | 1020 | 20 | 2 | 5.590 | 0.8
 | 0.0756 | 1.0
 | 1.864 | 1.3 | 0.1789 | 0.8
 | 0.63 |
| 2_16.1
2.24.1 | 0.00 | 21.0
510 | 5.06
211 | 3.45
88.4

 | 0.25 | 1129
1185 | 21
10 | 1115
1167 | 70
18 | -1
-2 | 5.220
4.956 | 2.0
1.0
 | 0.0768
0.0788 | 3.5
0.9
 | 2.026 | 4.1 | 0.1914
0.2018 | 2.0
1.0
 | 0.50 |
| 2_15.1 | 0.00 | 353 | 112 | 61.5

 | 0.33 | 1191 | 9.3 | 1185 | 22 | -1 | 4.926 | 0.9
 | 0.0795 | 1.1
 | 2.225 | 1.4 | 0.2030 | 0.9
 | 0.62 |
| 2_31.1
4 5.1 | 0.09
0.24 | 362
663 | 100
248 | 64.3
130

 | 0.29
0.39 | 1211
1321 | 9.6
8.5 | 1205
1391 | 22
16 | 0
5 | 4.840
4.396 | 0.9
0.7
 | 0.0803
0.0884 | 1.1
0.9
 | 2.288
2.772 | 1.4
1.1 | 0.2066
0.2274 | 0.9
0.7
 | 0.61
0.64 |
| 2_22.1 | 0.00 | 123 | 34.5 | 26.2

 | 0.29 | 1422 | 14 | 1400 | 29 | -2 | 4.053 | 1.1
 | 0.0888 | 1.5
 | 3.021 | 1.9 | 0.2467 | 1.1
 | 0.59 |
| 4_1.1
2_1.1 | 0.00 | 102
341 | 68.7
96.9 | 22.8
79.5

 | 0.69 | 1486
1548 | 12 | 1471
1541 | 23
17 | -1
0 | 3.856 | 0.9
 | 0.0922
0.0957 | 1.2
0.9
 | 3.296 3.581 | 1.5 | 0.2593 0.2715 | 0.9
 | 0.60 |
| 2_19.1 | 0.11 | 162 | 124 | 39.2

 | 0.79 | 1594 | 14 | 1606 | 25 | 1 | 3.563 | 1.0
 | 0.0990 | 1.3
 | 3.831 | 1.6 | 0.2806 | 1.0
 | 0.60 |
| 2_6.1
2_14.1 | 0.15 | 233
388 | 142 | 56.9
97.1

 | 0.63 | 1612
1648 | 15 | 1635 | 13
14 | -1 | 3.520 | 1.0
0.8
 | 0.1896 | 0.8
0.7
 | 4.039 | 1.3 | 0.2840 0.2912 | 1.0
0.8
 | 0.78 |
| 2_10.1 | 0.00 | 248 | 167 | 63.0

 | 0.70 | 1671 | 14 | 1663 | 20 | -1 | 3.379 | 0.9
 | 0.1021 | 1.1
 | 4.167 | 1.4 | 0.2959 | 0.9
 | 0.65 |
| 2_7.1
2_30.1 | 0.00 | 235
291 | 89.1 | 60.9
77.8

 | 0.53 | 1698 | 13 | 1682 | 17 | -1 | 3.319 | 0.9
 | 0.1032 | 0.9
 | 4.287
4.543 | 1.3 | 0.3013 | 0.9
 | 0.70 |
| 2_13.1 | 0.10 | 211 | 87.8 | 58.5

 | 0.43 | 1804 | 14 | 1819 | 18 | 1 | 3.096 | 0.9
 | 0.1112 | 1.0
 | 4.951 | 1.3 | 0.3229 | 0.9
 | 0.68 |
| 2_17.1
2_28.1 | 0.02 | 425
86.6 | 57.8 | 25.6

 | 0.23 | 1908 | 20 | 1911 | 26
25 | 0 | 2.984 | 1.0
 | 0.1262 | 1.5
 | 5.560 | 1.8 | 0.3331 | 1.0
 | 0.54 |
| 2_27.1 | 0.00 | 408 | 143 | 121

 | 0.36 | 1913 | 14 | 2279 | 11 | 19 | 2.895 | 0.8
 | 0.1442 | 0.6
 | 6.871 | 1.0 | 0.3455 | 0.8
 | 0.81 |
| 2_20.1
2_29.1 | 0.00 | 80.5 | 77.5 | 25.1

 | 1.00 | 1994 | 20 | 1921 | 23 | 0 | 2.759 | 1.2
 | 0.1220 | 1.3
 | 6.100 | 1.7 | 0.3624 | 1.2
 | 0.68 |
| 2_3.1
2_4_1 | | 801
51.2 | 37.0
57.2 | 294
19.6

 | 0.05 | 2293
2382 | 18
32 | 2727
2350 | 7.2
39 | 19
-1 | 2.341 | 0.9
 | 0.1882 | 0.4
 | 11.09
9.270 | 1.0 | 0.4273 | 0.9
 | 0.91 |
| 2_2.1 | 0.00 | 186 | 209 | 78.5

 | 1.16 | 2571 | 20 | 2691 | 11 | 5 | 2.041 | 1.0
 | 0.1842 | 0.7
 | 12.45 | 1.2 | 0.4901 | 1.0
 | 0.81 |
| 2_32.1
2_25.1 | 0.05 | 251
403 | 15.6
69.4 | 110
183

 | 0.06 | 2667
2732 | 23
18 | 2747
2708 | 9.9
8.5 | 3 | 1.952
1.894 | 1.0
0.8
 | 0.1906 | 0.6
0.5
 | 13.46
13.54 | 1.2 | 0.5123 | 1.0
0.8
 | 0.87
0.85 |
| 2 11.1 | 0.00 | 202 | 676 | 96.1

 | 0.03 | 2834 | 21 | 2994 | 9.3 | 6 | 1.811 | 0.9
 | 0.2210 | 0.6
 | 16.80 | 1.1 | 0.5522 | 0.9
 | 0.85 |
| | 0.00 | 202 | 0.70 | 50.1

 | 0.05 | 20.54 | 21 | 2004 | | | | 0.7
 | 0.2219 | 0.0
 | 10.89 | 1.1 | 0.3322 |
 | 0.05 |
| 2_8.1
Sandstone f | 0.00
ragments | 112
from the M | 40.3
Aendeleev | 58.6
Rise, sam

 | 0.03
0.37
ple AF05-15 | 3061
(78°58'N, | 29
173°56'W) | 3082 | 11 | 1 | 1.645 | 1.2
 | 0.2344 | 0.0
 | 19.64 | 1.1 | 0.6078 | 1.2
 | 0.86 |
| 2_8.1
Sandstone f
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31.1 | 0.00
ragments
0.36
0.00 | 112
from the M
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495 | 40.3
40.3
1endeleev
784
160 | 58.6
Rise, sam
34.0
28.7

 | 0.03
0.37
ple AF05-15
0.86
0.33 | 3061
(78°58'N,
265
421 | 29
173°56'W)
10
4.2 | 533
390 | 11
97
45 | 1
101
-7 | 1.645
23.82
14.82 | 3.9
1.0
 | 0.2219
0.2344
0.0581
0.0545 | 0.0
0.7
4.4
2.0
 | 0.336
0.507 | 1.1
1.4
5.9
2.3 | 0.0420 | 1.2
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 | 0.86 |
| 2_8.1
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<u>Aendeleev</u>
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0.0712 | 0.0
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0.1787 | 3.9
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Table 1. Continued.

Analysis	%	ppm	ppm	ppm		(1) 206	Ph/ ²³⁸ U	(1) ²⁰⁷ Ph	/ ²⁰⁶ Pb	%	(1)		(1)		(1)		(1)		err
number	²⁰⁶ Pb _c	U	Th	²⁰⁶ Pb*	²³² Th/ ²³⁸ U	Age (Ma	a) ± abs	Age (Ma)	± abs	Dis.	$^{238}\mathrm{U/}^{206}\mathrm{Pb}^{*}$	±%	²⁰⁷ Pb ^{*/206} Pb [*]	±%	207Pb*/235U	±%	206Pb*/238U	±%	corr
5.1	0.40	49.6	25.7	9.89	0.54	1341	13	1343	51	0	4.323	1.1	0.0862	2.6	2.750	2.8	0.2313	1.1	0.38
9.1	0.07	138	45.5	28.1	0.34	1373	8.4	1390	20	1	4.213	0.7	0.0883	1.0	2.892	1.2	0.2374	0.7	0.54
19.1	0.12	451	162	94.1	0.37	1400	4.4	1460	11	4	4.123	0.3	0.0916	0.6	3.064	0.7	0.2426	0.3	0.51
28.1	0.03	247	107	51.6	0.45	1400	5.7	1444	15	3	4.121	0.5	0.0909	0.8	3.039	0.9	0.2426	0.5	0.50
23.1	0.10	188	74.8	39.6	0.41	1412	6.5	1455	21	3	4.083	0.5	0.0914	1.1	3.086	1.2	0.2449	0.5	0.42
27.1	0.04	96.7	63.1	20.7	0.67	1432	10	1442	22	1	4.021	0.8	0.0908	1.1	3.113	1.4	0.2487	0.8	0.56
7.1		186	88.3	40.1	0.49	1443	6.4	1440	16	0	3.985	0.5	0.0907	0.8	3.138	1.0	0.2510	0.5	0.51
4.1		126	61.3	27.7	0.50	1476	8.8	1477	19	0	3.886	0.7	0.0925	1.0	3.281	1.2	0.2573	0.7	0.56
20.1	0.04	164	60.9	37.7	0.38	1524	8.3	1667	16	9	3.751	0.6	0.1023	0.9	3.762	1.1	0.2666	0.6	0.58
14.1	0.35	53.1	38.3	12.9	0.74	1596	12	1590	32	0	3.561	0.9	0.0982	1.7	3.801	1.9	0.2808	0.9	0.45
13.1	0.05	115	112	28.2	1.01	1623	10	1658	18	2	3.494	0.7	0.1018	1.0	4.018	1.2	0.2862	0.7	0.59
21.1	0.01	511	227	126	0.46	1624	4.7	1733	8.1	7	3.490	0.3	0.1061	0.4	4.190	0.5	0.2866	0.3	0.59
10.1	0.06	149	75.3	37.8	0.52	1660	7.7	1684	16	1	3.404	0.5	0.1033	0.9	4.183	1.0	0.2938	0.5	0.52
25.1	0.00	149	45.8	38.5	0.32	1697	9.1	1698	15	0	3.320	0.6	0.1041	0.8	4.323	1.0	0.3012	0.6	0.61
16.1	0.00	302	104	84.2	0.36	1813	6.1	1871	9.2	3	3.080	0.4	0.1145	0.5	5.124	0.6	0.3247	0.4	0.60
3.1	0.08	96.2	52.5	28.0	0.56	1877	10	1917	16	2	2.958	0.6	0.1174	0.9	5.472	1.1	0.3381	0.6	0.57
29.1	0.00	258	255	76.3	1.02	1911	7.7	1906	10	0	2.899	0.5	0.1167	0.5	5.551	0.7	0.3450	0.5	0.65
12.1		83.7	67.7	37.1	0.84	2685	14	2681	15	0	1.936	0.6	0.1830	0.9	13.04	1.1	0.5166	0.6	0.57
30.1	0.00	40.1	17.7	18.4	0.46	2757	23	2751	18	0	1.873	1.0	0.1910	1.1	14.06	1.5	0.5338	1.0	0.69

Note: Dots in the first column separate numbers of analyzed grains from numbers of shots. Errors marked "+/- abs" and "+/- %" are within 1 sigma. Pbc and Pb* indicate common and radiogenic lead, respectively.

Columns (1) designate radiogenic Pb calculated using measured ²⁰⁴Pb, Relative discordancy (% Dis) is calculated as 100 x [[age[207/206] / [age[206/238)] – 1]. Error correlation (err corr) is the correlation of errors for Pb/U isotope ratio.

Analysis number	²⁰⁶ Pb/ ²³⁸ U	± 1s, abs	²⁰⁷ Pb/ ²⁰⁶ Pb	± 1s, abs	²⁰⁶ Pb/ ²³⁸ U Age (Ma)	± 1s, abs	²⁰⁷ Pb/ ²⁰⁶ Pb Age (Ma)	± 1s, abs
1.1	0.2862	0.0032	0.0955	0.0013	1636	21	1533	49
2.1	0.0401	0.0016	0.0630	0.0008	196	17	755	40
3.1	0.1685	0.0066	0.0767	0.0025	878	200	1066	220
4.1	0.1066	0.0011	0.2566	0.0048	662	22	3216	98
5.1	0.1323	0.0062	0.1383	0.0011	568	12	2132	5.3
6.1	0.0618	0.0004	0.0599	0.0011	386	5.6	526	13
7.1	0.3051	0.0026	0.1066	0.0011	1703	33	1709	56
8.1	0.1891	0.0029	0.0793	0.0009	1091	18	1143	21
9.1	0.3257	0.0017	0.1078	0.0002	1832	32	1770	6.8
10.1	0.0302	0.0018	0.0687	0.0012	170	20	818	100
11.1	0.2323	0.0015	0.0794	0.0013	1352	20	1197	35
12.1	0.3223	0.0059	0.1018	0.0008	1840	85	1644	52
13.1	0.3151	0.0013	0.1056	0.0001	1758	28	1729	2.4
14.1	0.1613	0.0096	0.1062	0.0003	933	150	1747	18
15.1	0.1229	0.0057	0.1184	0.0002	689	100	1942	14
16.1	0.3774	0.0061	0.1499	0.0003	2001	88	2334	16
17.1	0.2251	0.0106	0.0958	0.0007	1342	170	1513	38
18.1	0.3328	0.0248	0.1963	0.0041	2001	130	2731	31
19.1	0.0768	0.0010	0.1262	0.0010	479	9.5	2029	9.9
20.1	0.2478	0.0089	0.1084	0.0007	1646	120	1829	22
21.1	0.1918	0.0042	0.0968	0.0044	1169	71	1383	270
22.1	0.2841	0.0014	0.1123	0.0008	1615	29	1820	24
23.1	0.2901	0.0014	0.1061	0.0002	1653	28	1731	13
24.1	0.4707	0.0169	0.1935	0.0008	2376	100	2787	14
25.1	0.1066	0.0011	0.2566	0.0048	1303	76	1652	50
26.1	0.1066	0.0011	0.2566	0.0048	1240	100	1768	14
27.1	0.1176	0.0018	0.1209	0.0003	691	17	1961	5.5
28.1	0 2489	0.0063	0.1052	0.0004	1425	58	1715	71
29.1	0.1671	0.0102	0.1194	0.0005	800	30	1953	3.0
30.1	0.4831	0.0035	0.1630	0.0003	2518	56	2502	13
31.1	0.4215	0.0029	0.1548	0.0009	2312	20	2406	37
32.1	0.2801	0.0058	0.1095	0.0002	1587	130	1790	13
33.1	0.1863	0.0060	0.0976	0.0032	1141	97	1501	140
34.1	0.3239	0.0032	0.1964	0.0073	1800	28	2774	90
35.1	0.2673	0.0130	0.1079	0.0010	1379	170	1750	53
36.1	0.0904	0.0015	nd	nd	536	37	nd	nd
37.1	0.2980	0.0031	0.1136	0.0003	1631	26	1861	2.2
38.1	0.2874	0.0010	0.1046	0.0010	1626	23	1725	23
39.1	0.3542	0.0027	0.1150	0.0003	1929	30	1880	6.5
40.1	0.1796	0.0011	0.0761	0.0004	1058	37	1087	35
41.1	0.1908	0.0020	0.0810	0.0022	1126	31	1158	170
42.1	0.2816	0.0138	0.0984	0.0010	1783	130	1655	26
43.1	0.2427	0.0098	0.1129	0.0004	1531	160	1823	25
44.1	0.1549	0.0990	-0.1858	0.4786	583	190	2104	18
45.1	0.3415	0.0036	0.1092	0.0004	1934	44	1770	15
46.1	0.2253	0.0018	0.0972	0.0005	1316	30	1578	56
47.1	0.1134	0.0038	0.0894	0.0007	732	35	1400	21
48.1	0.1867	0.0037	0.1197	0.0002	1161	18	1948	2.7
49.1	0.3046	0.0029	0.1054	0.0003	1740	25	1721	3.1
50.1	0.2865	0.0057	0.1197	0.0006	1614	27	1969	4.3

Table 2. U-Pb analytical data obtained by laser ablation coupled with MC-ICPMS-HR NEPTUNE (Thermo TM). Sandstone fragment from the Lomonosov Ridge, sample ALR07-18 (82°30'N, 140°E)

Note: nd - corresponds to extremely low-Pb content for apropriate estimation of Pb/Pb and U/Pb ratios.

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